



RGF Chantier Pyrénées

**Gravity anomalies, lateral variation of flexural rigidity, and
flexure of a rifted continental margin: the Aquitaine Basin**

Paul ANGRAND¹, Mary FORD¹, Anthony WATTS²

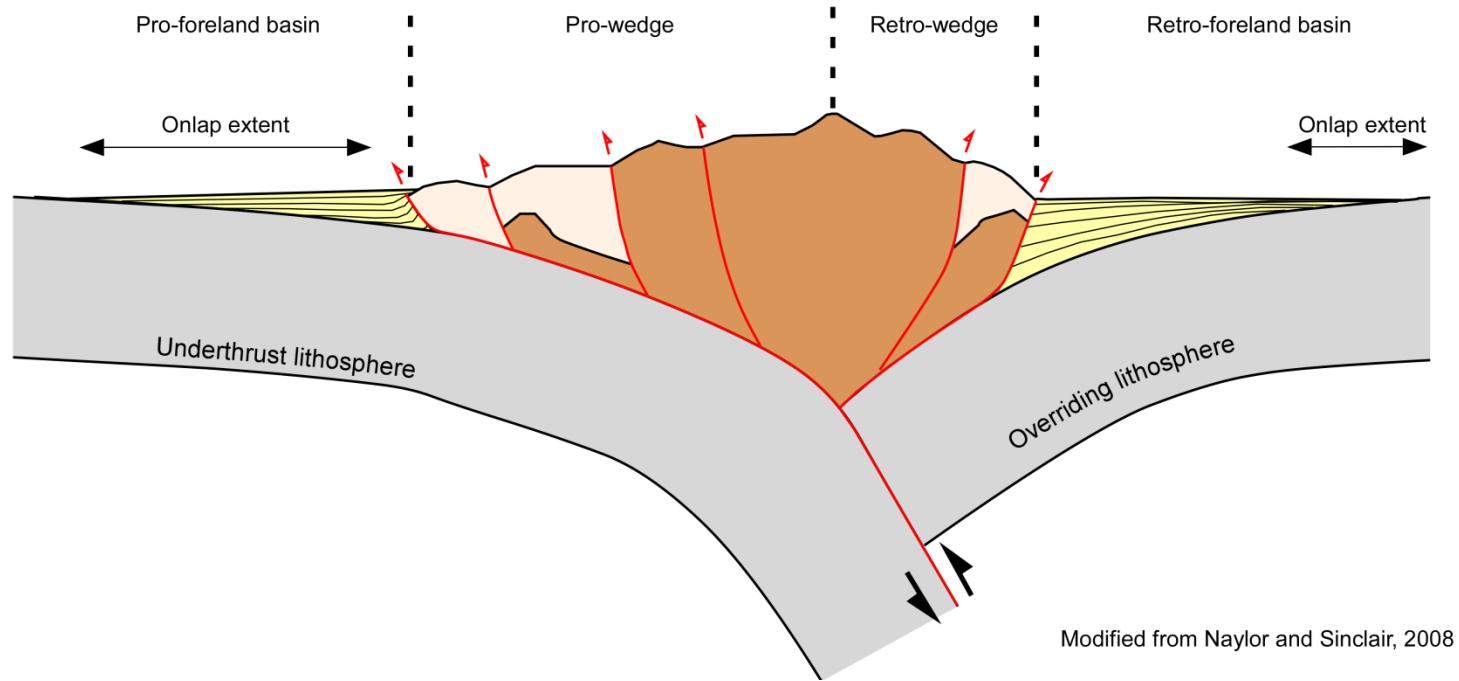
1: CRPG-CNRS Université de Lorraine, Nancy; 2: Oxford University, Oxford

Journées RGF 22-23 mars 2017

Chantier Pyrénées

Background

Foreland basins

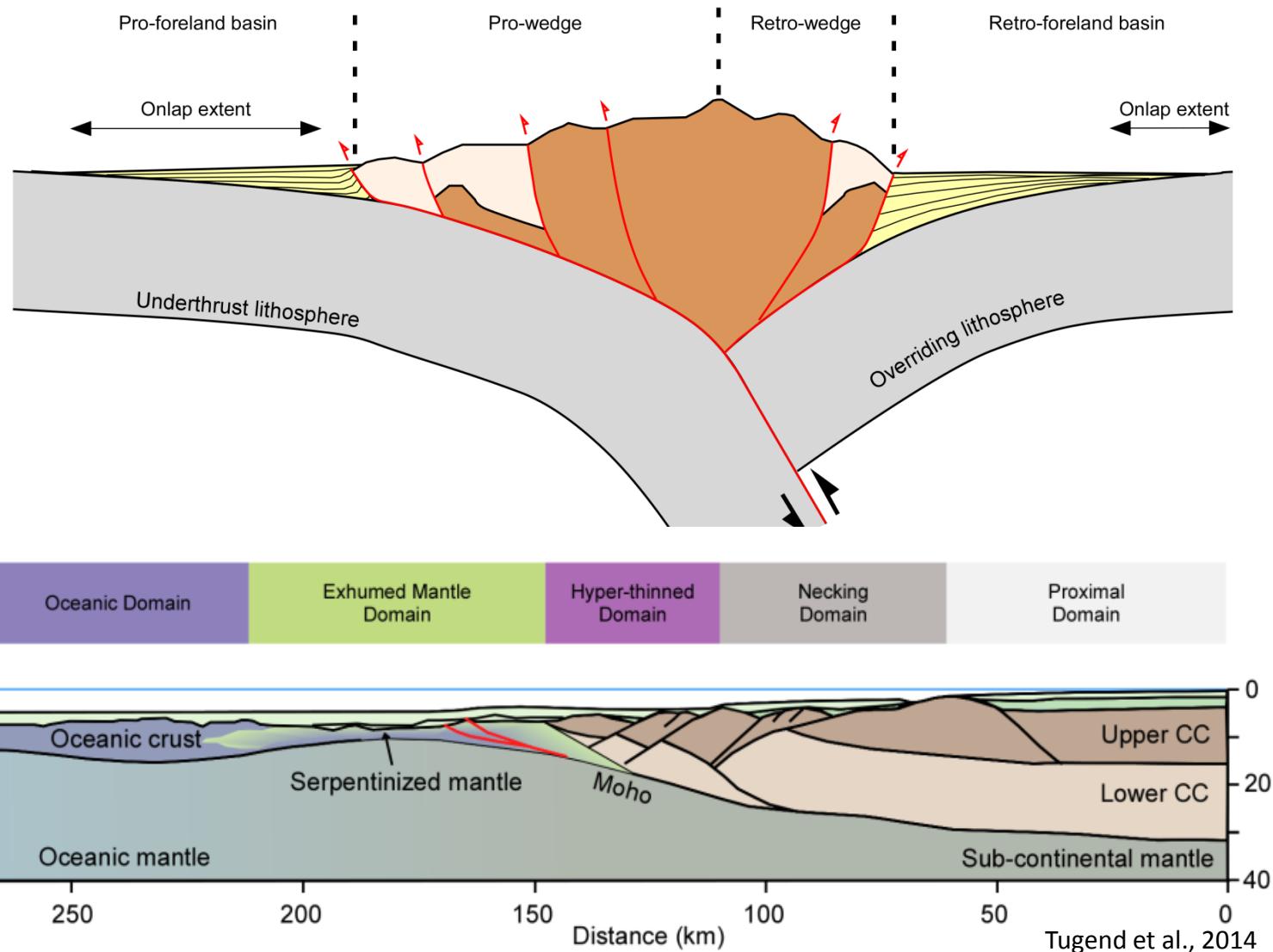


Flexural foreland basins that develop due to the load of the orogen

Geometry depends on the rigidity of the plate (T_e)

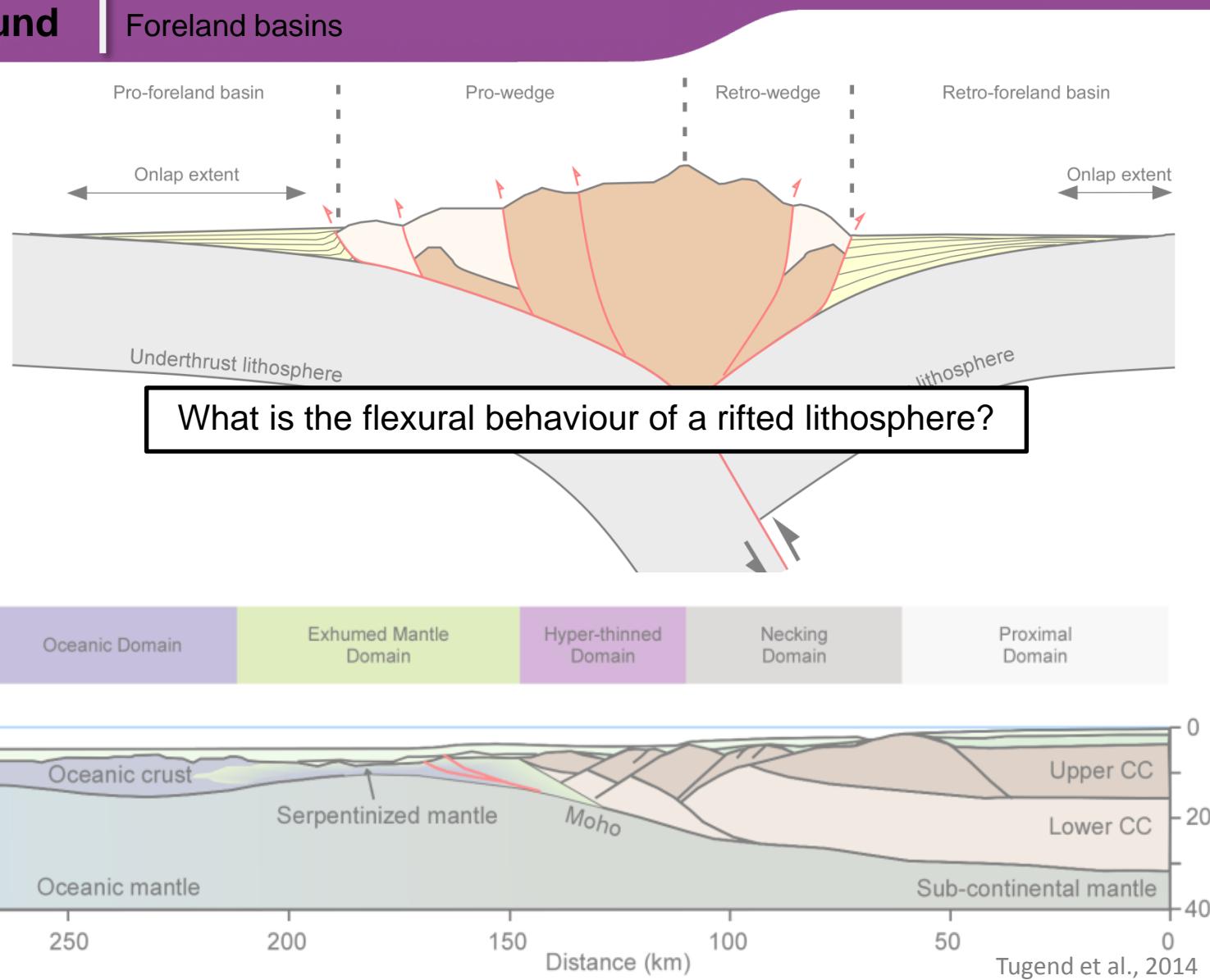
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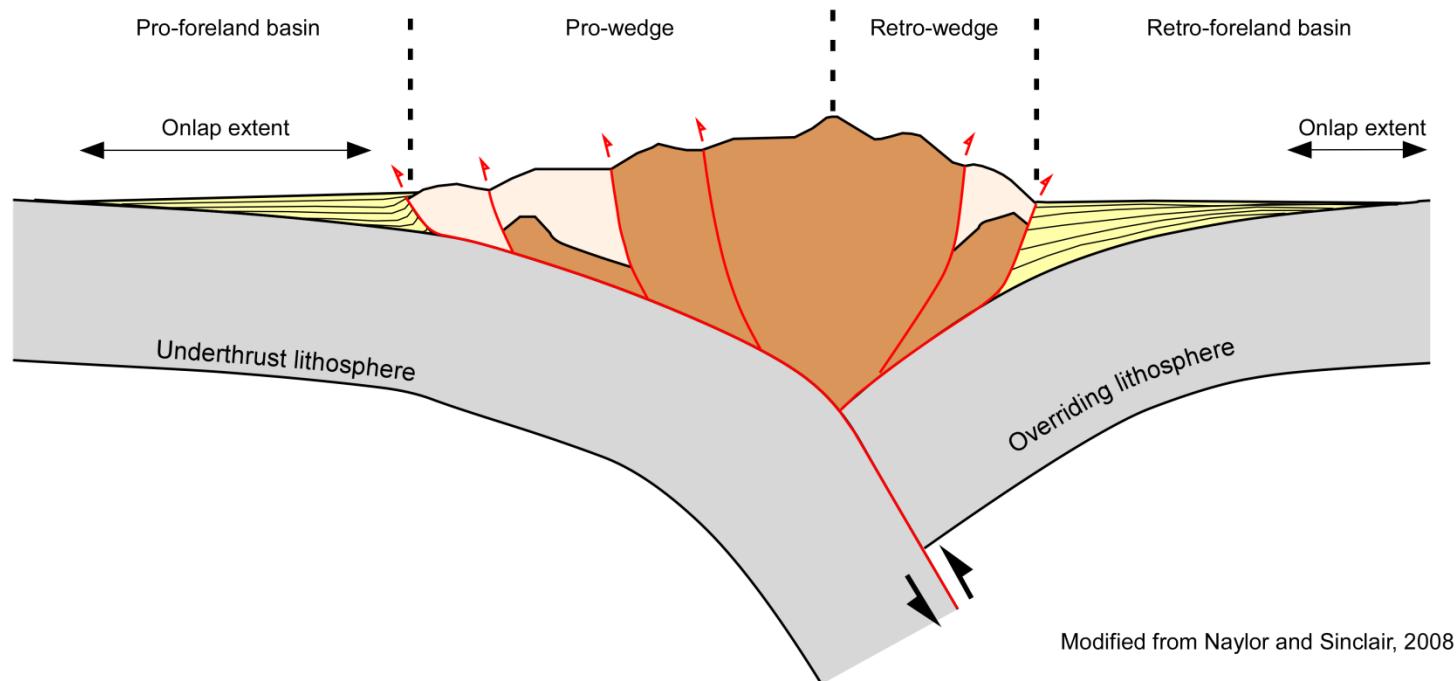
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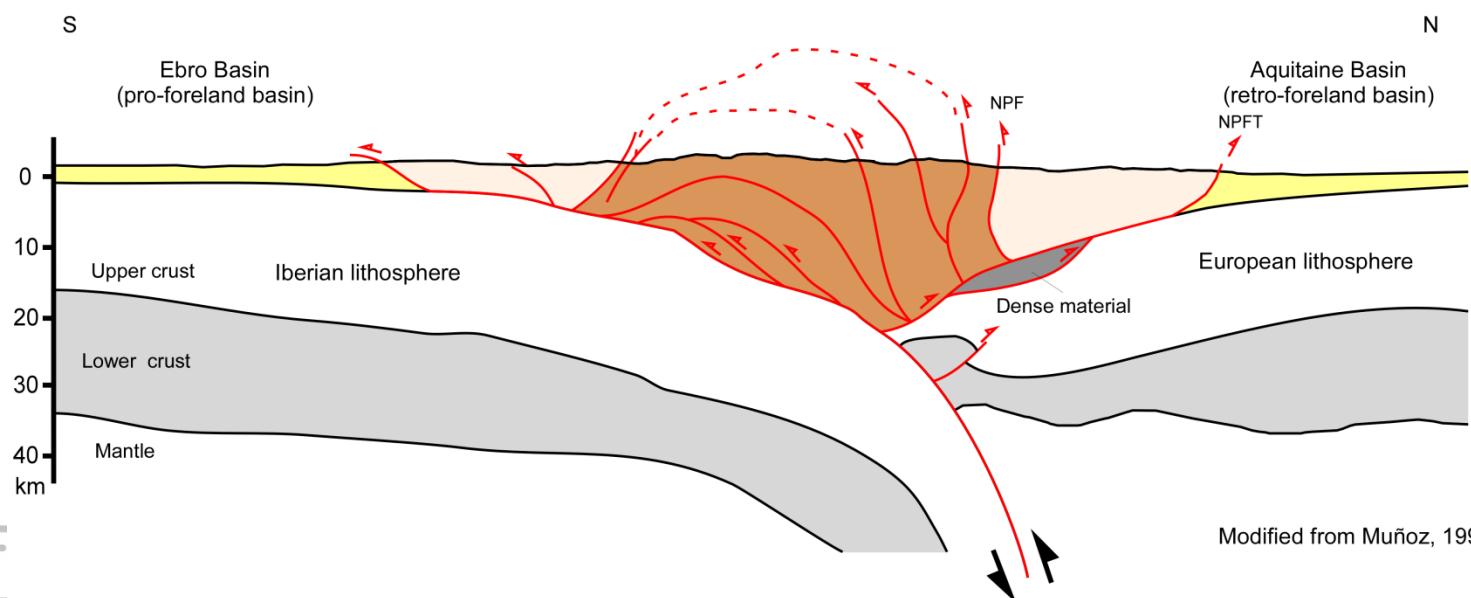


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Foreland basins



Modified from Naylor and Sinclair, 2008

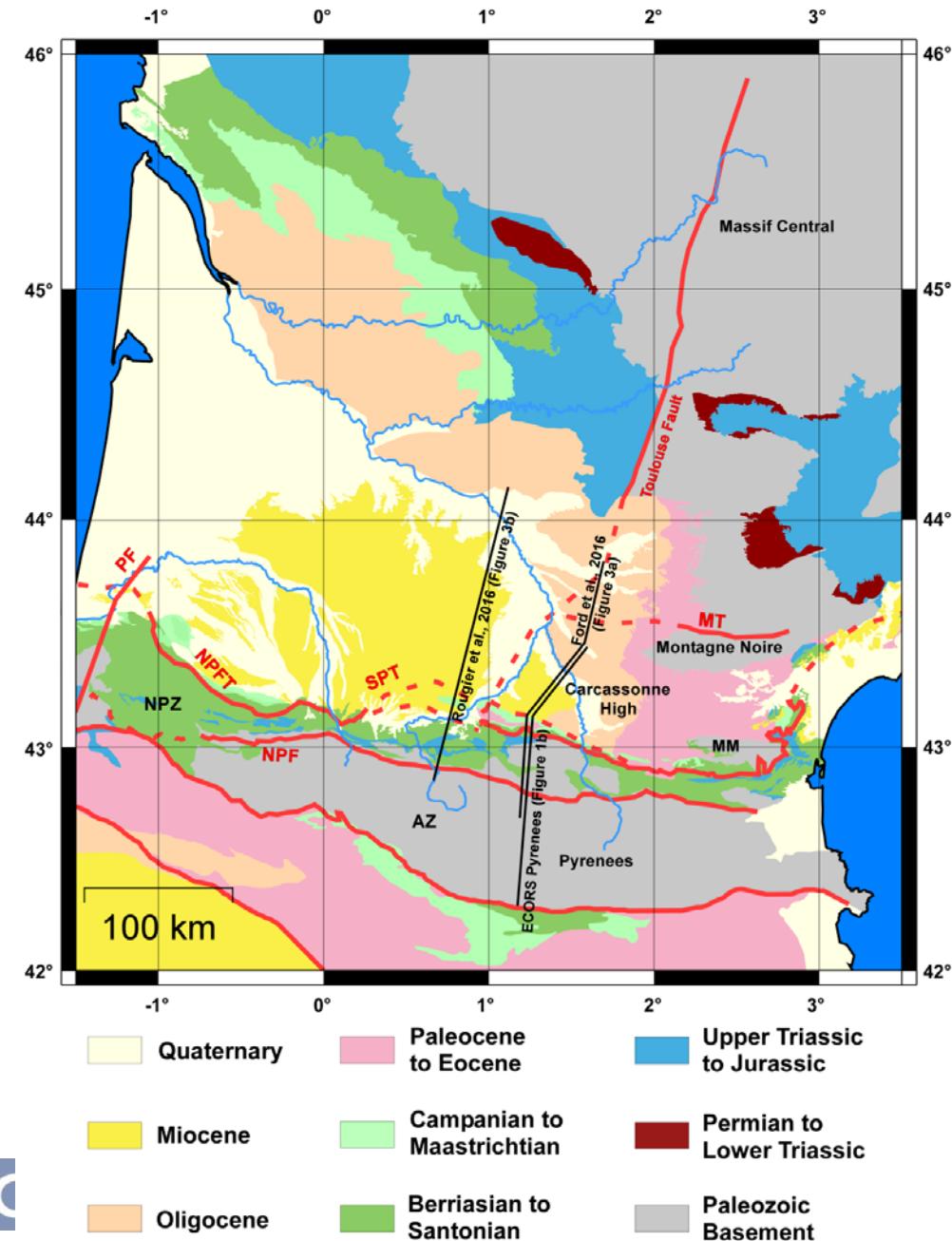


Modified from Muñoz, 1992

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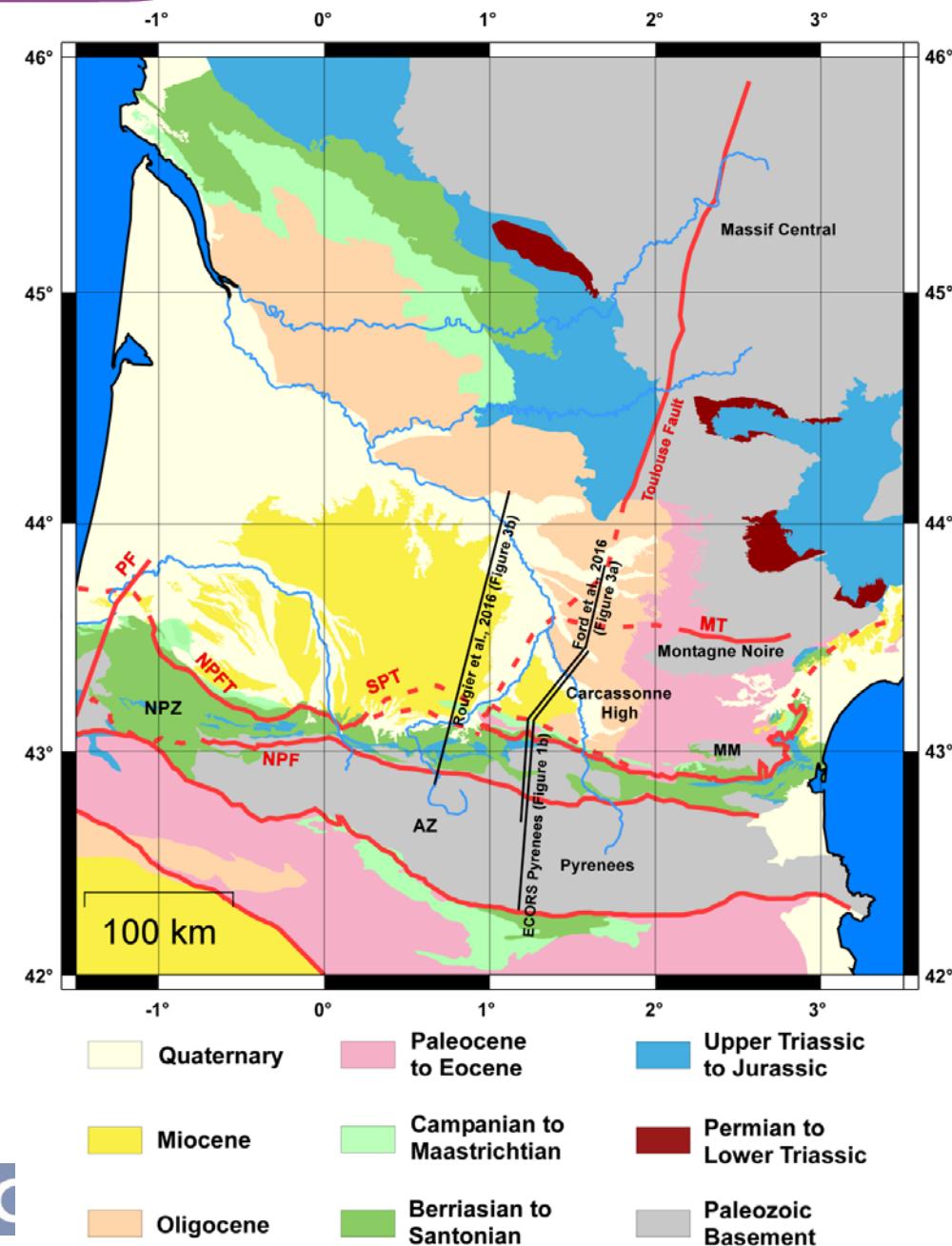
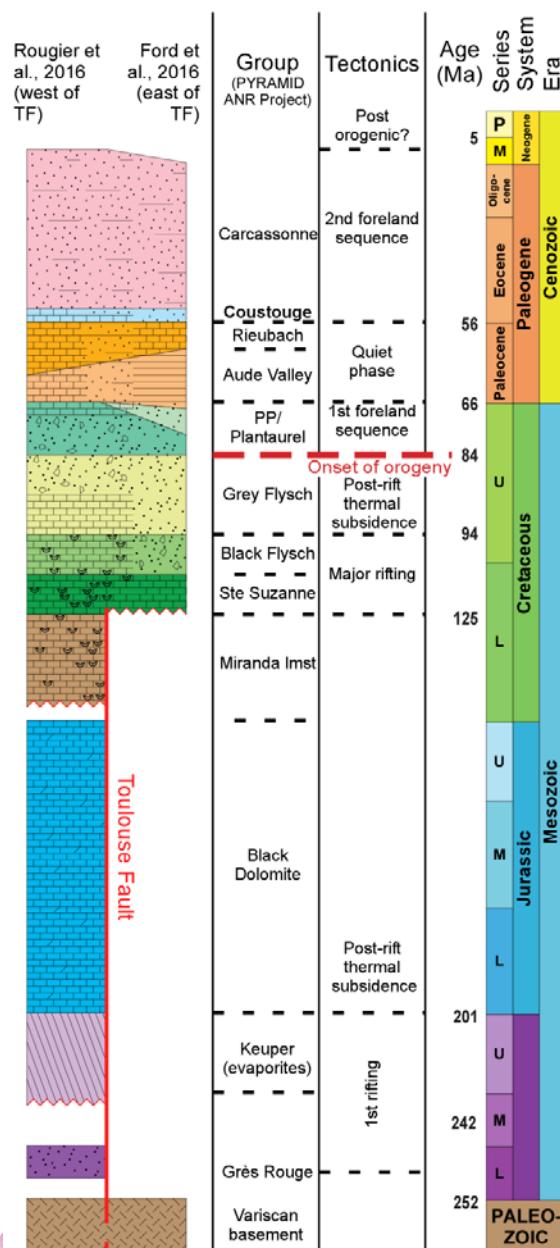
Geology of the Aquitaine Basin

- Pyrenean retro-foreland basin
- Syn-orogenic: Campanian to Miocene
- Asymmetric
- Structural inheritance (Apto-Cenomanian rifting: opening of the Bay of Biscay)
- Inherited faults
- Also post-orogenic events:
 - Opening of the Gulf of Lion
 - Pliocene uplift of the Massif Central



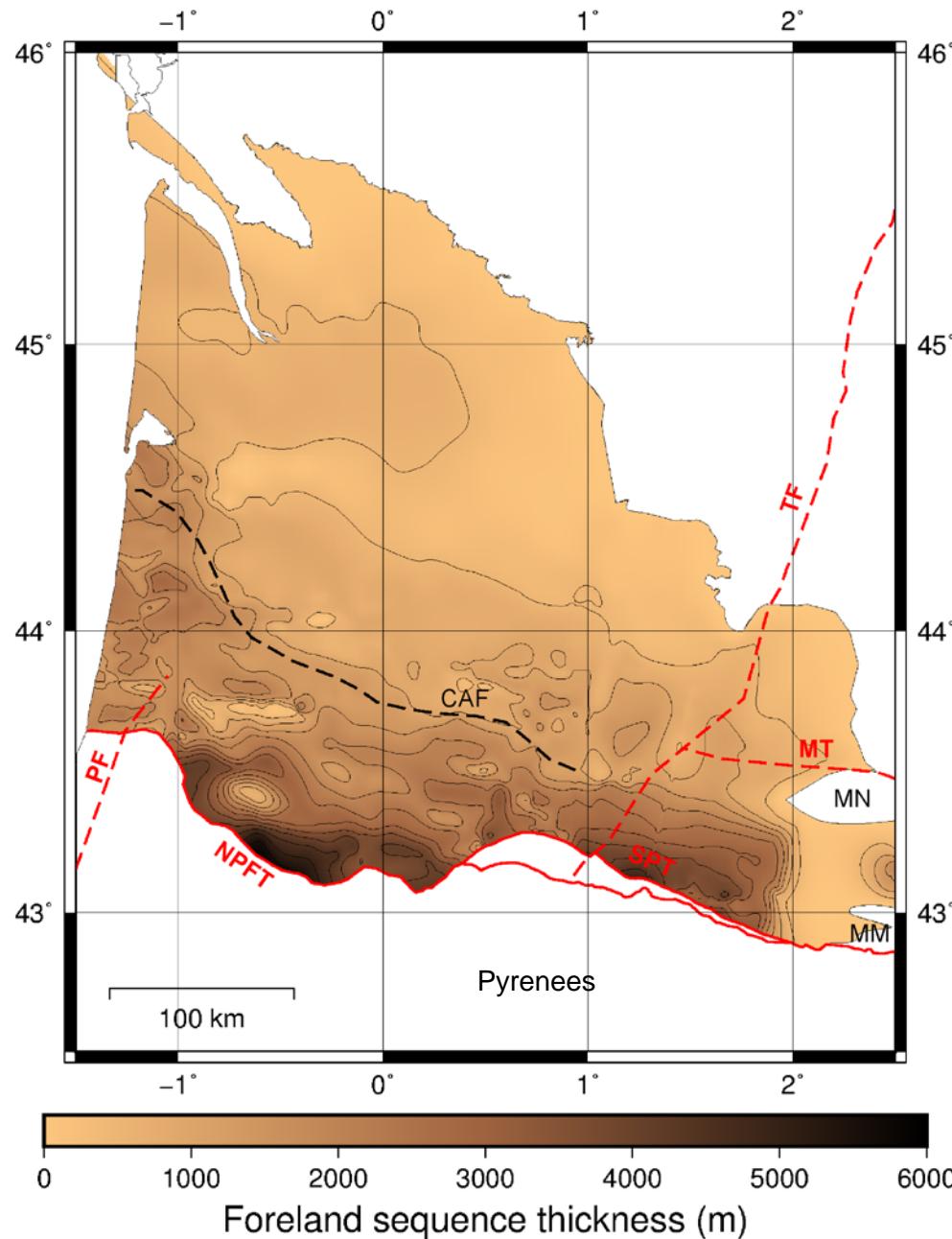
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Geology of the Aquitaine Basin



The Aquitaine foreland basin

Foreland distribution



- Campanian to Miocene
- Data:
 - Isopach maps (BRGM et al., 1974)
 - PYRAMID cross-sections (Ford et al., 2016; Rougier et al., 2016; Espurt et al., in prep., Ford et al., in prep.; Grool et al., in prep.)
- Thickening southward (up to 5.5 km thick)
- Flexure of ~150 km wide and bulge of 20-50 km wide

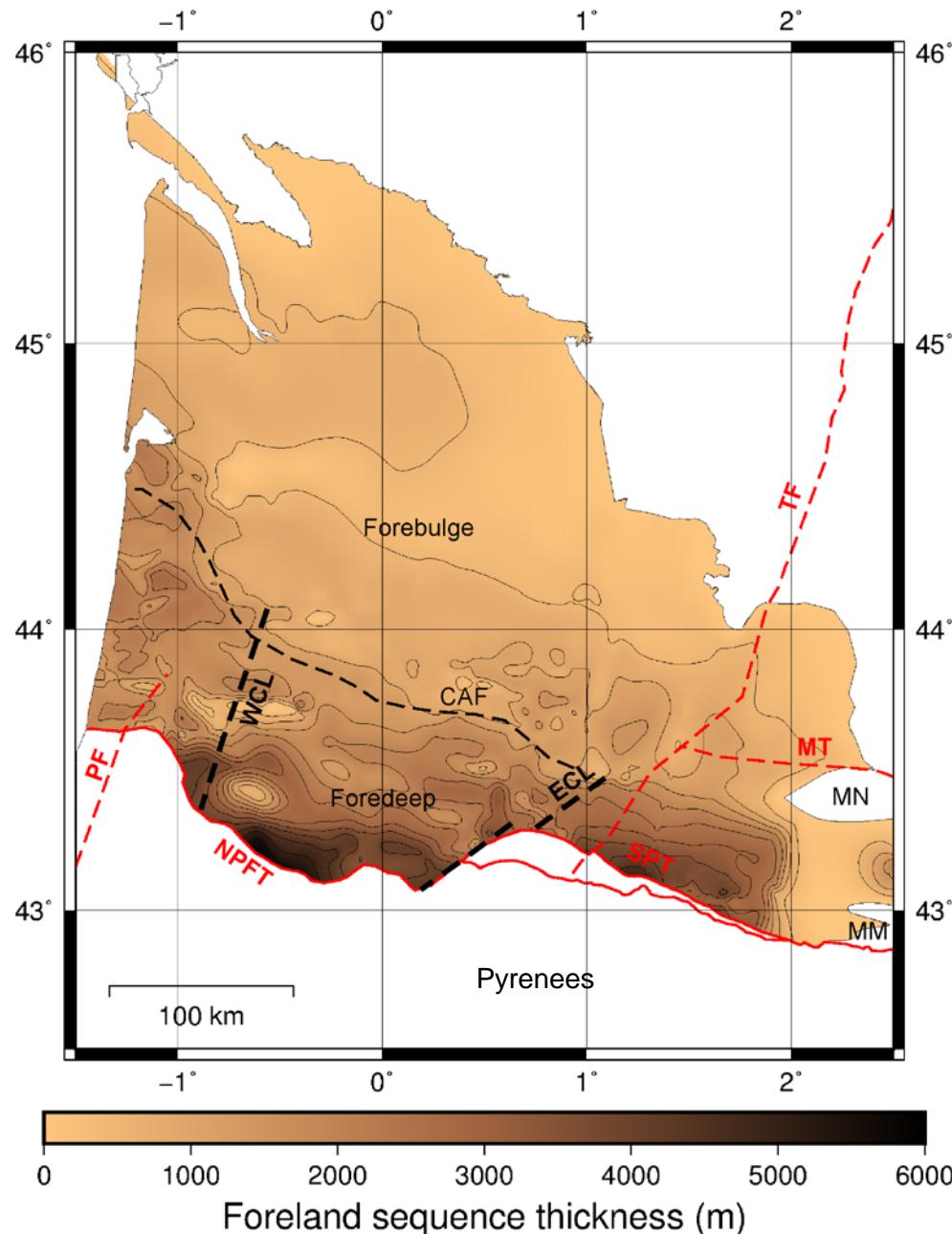
Angrand et al., in prep.



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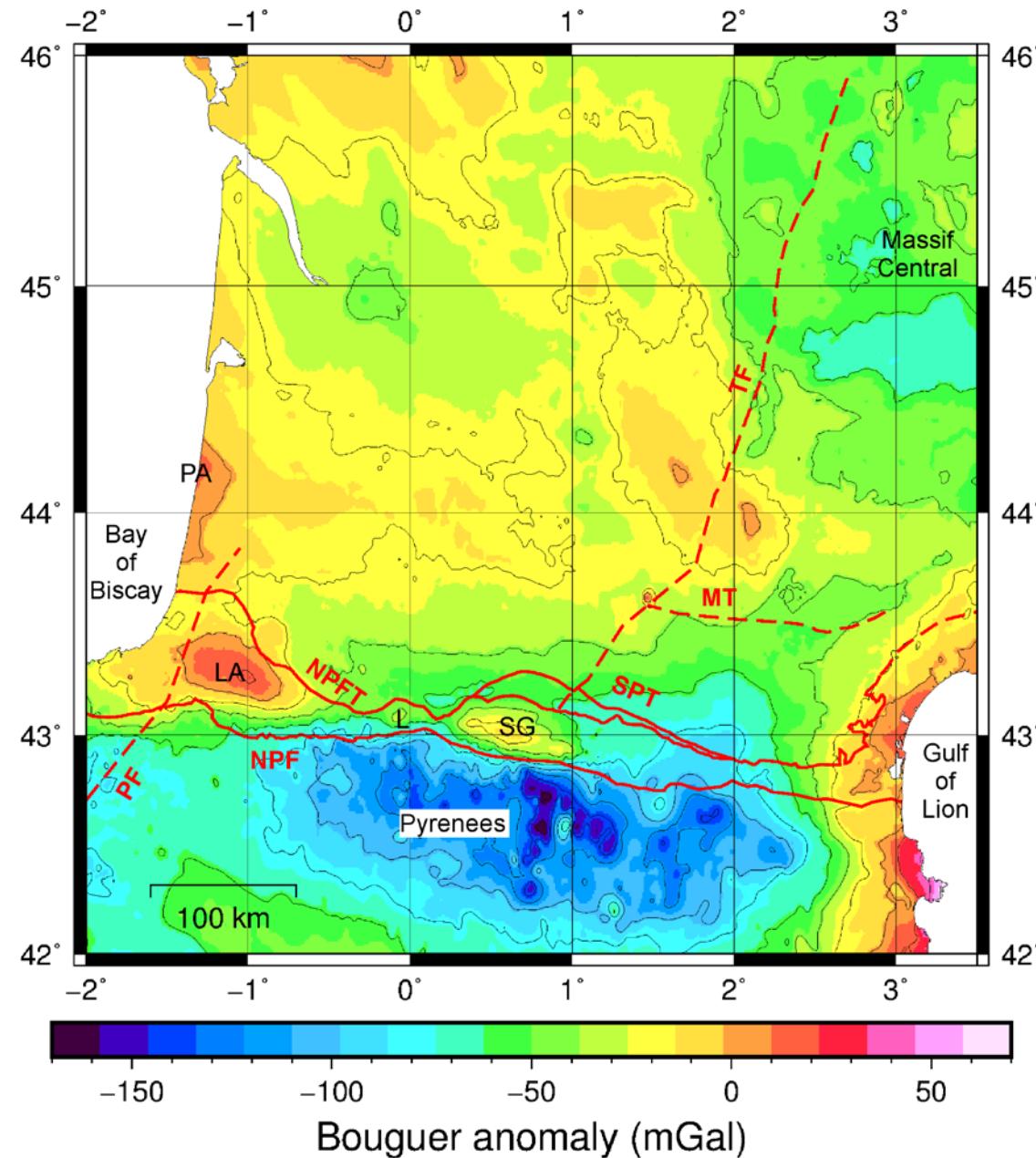
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The Aquitaine foreland basin

Bouguer anomaly

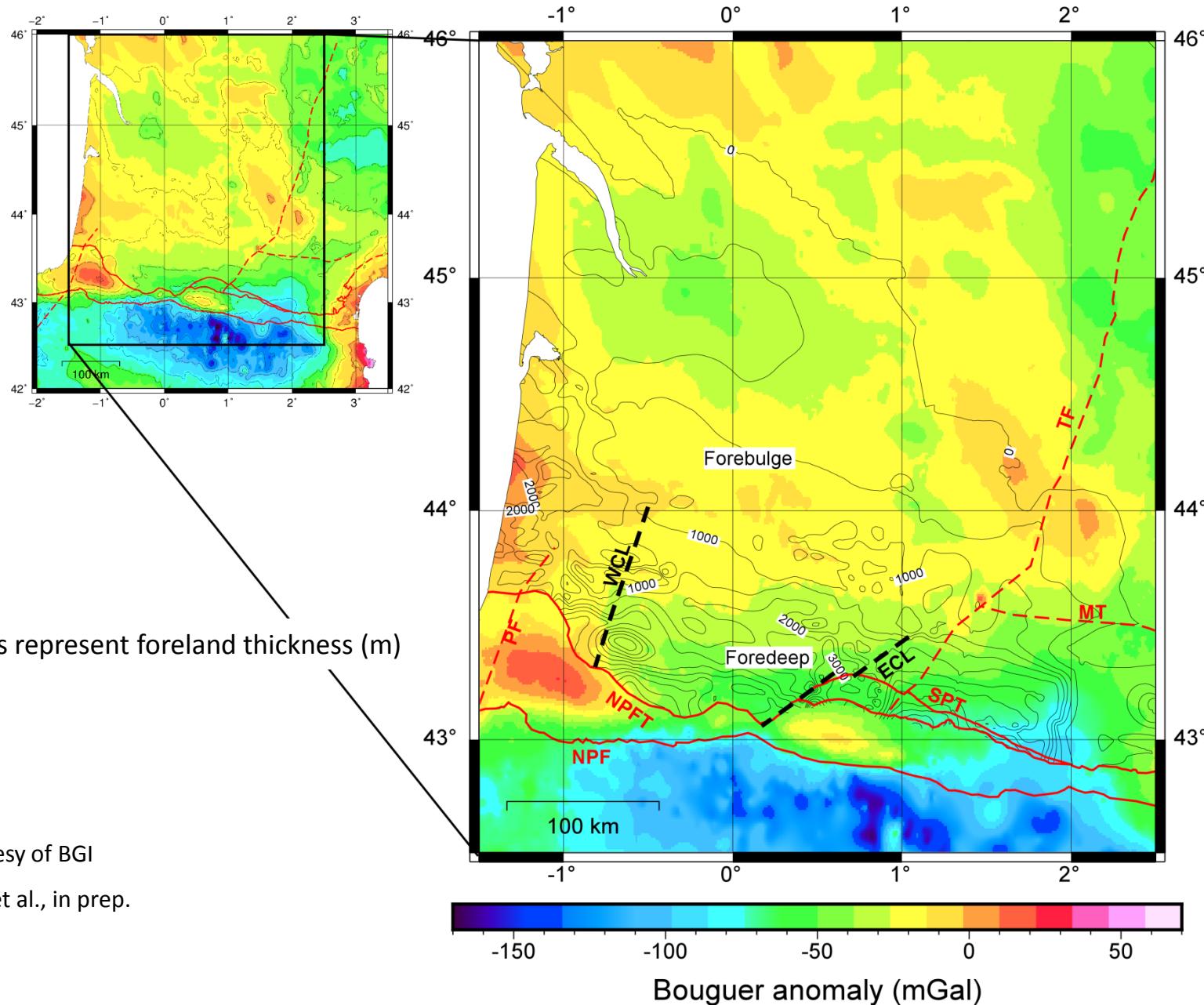


Courtesy of BGI

Angrand et al., in prep.

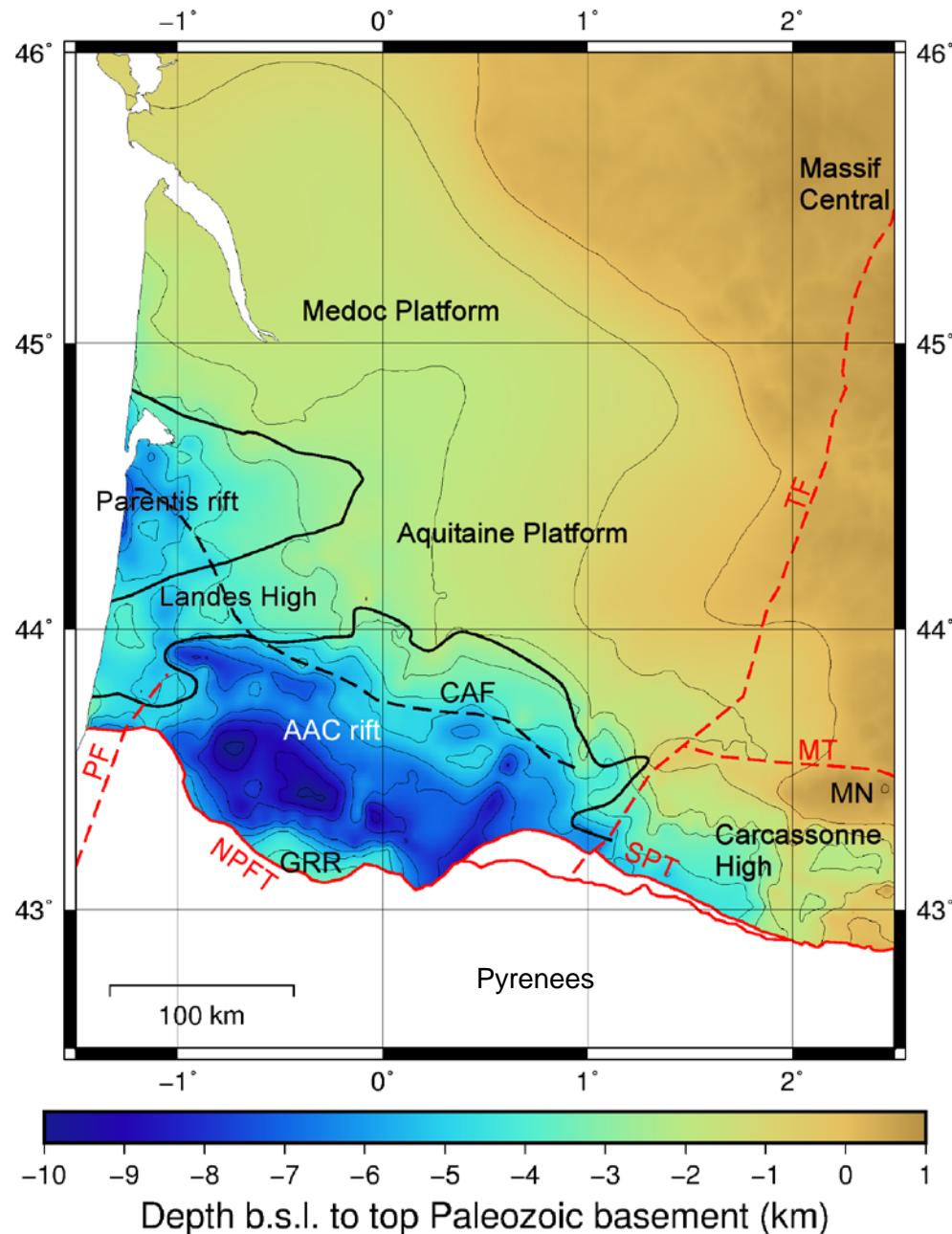
The Aquitaine foreland basin

Bouguer anomaly



Crustal template

Depth of top Paleozoic basement



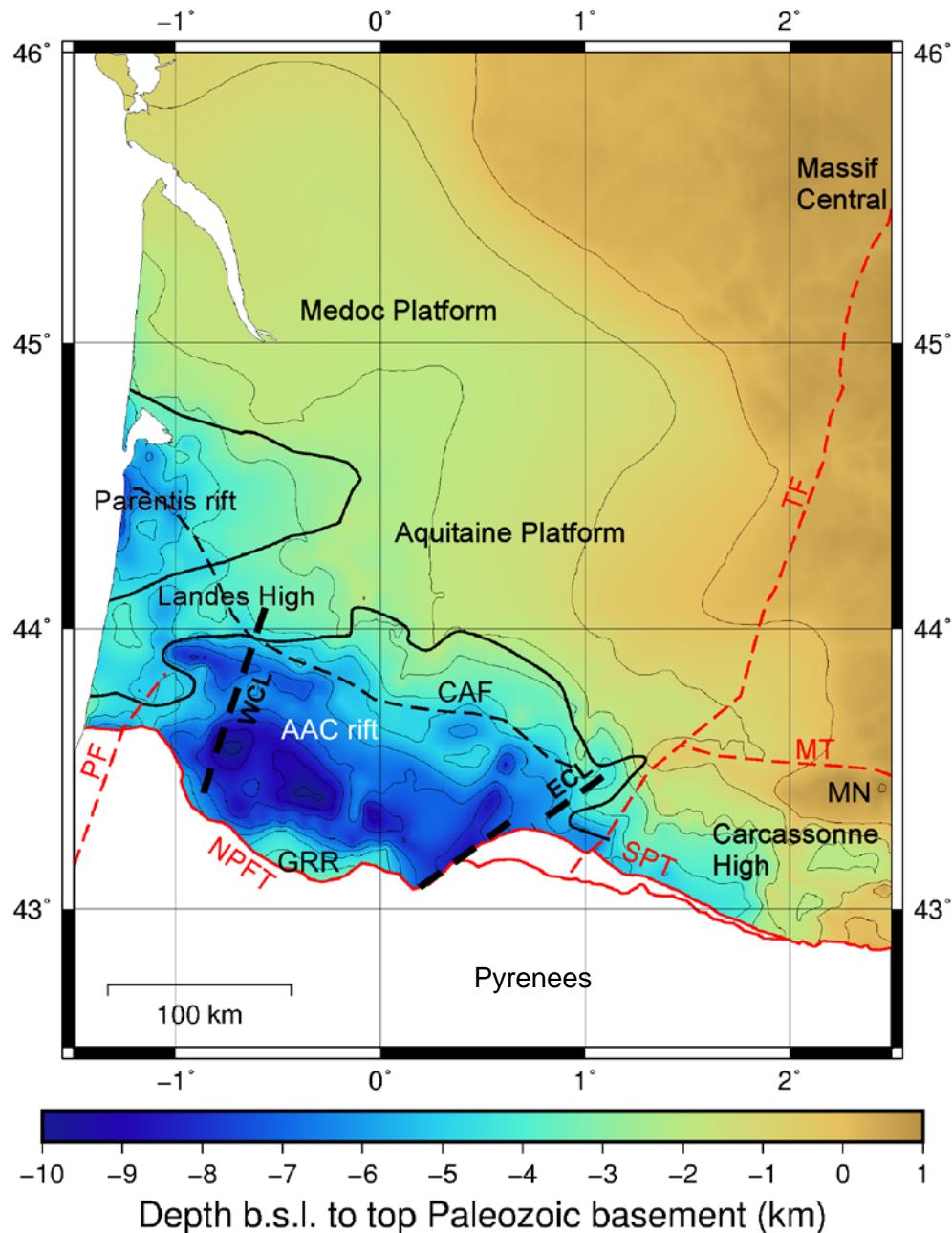
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- Deepens southward (down to almost 10 km in the Adour sub-basin)

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Crustal template

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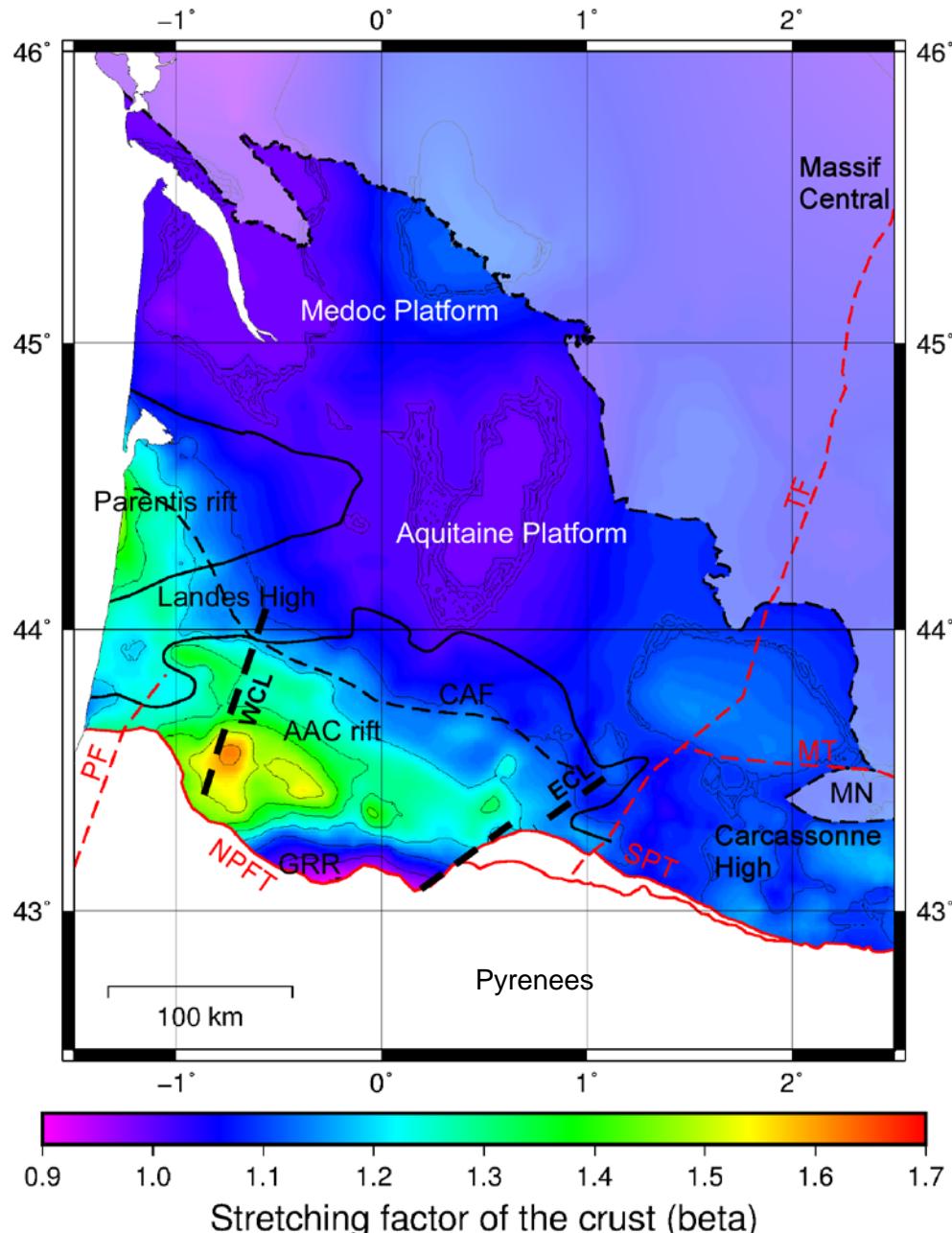
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Crustal template

Stretching of the plate



- Top of the crust: depth to top Paleozoic basement
- Base of the crust: depth to Moho, compilation of seismic data (Navarro et al., 2006)
- Initial crustal thickness: 33 km (from the platform)

$$\beta = \frac{\text{initial crustal thickness}}{\text{final crustal thickness}}$$

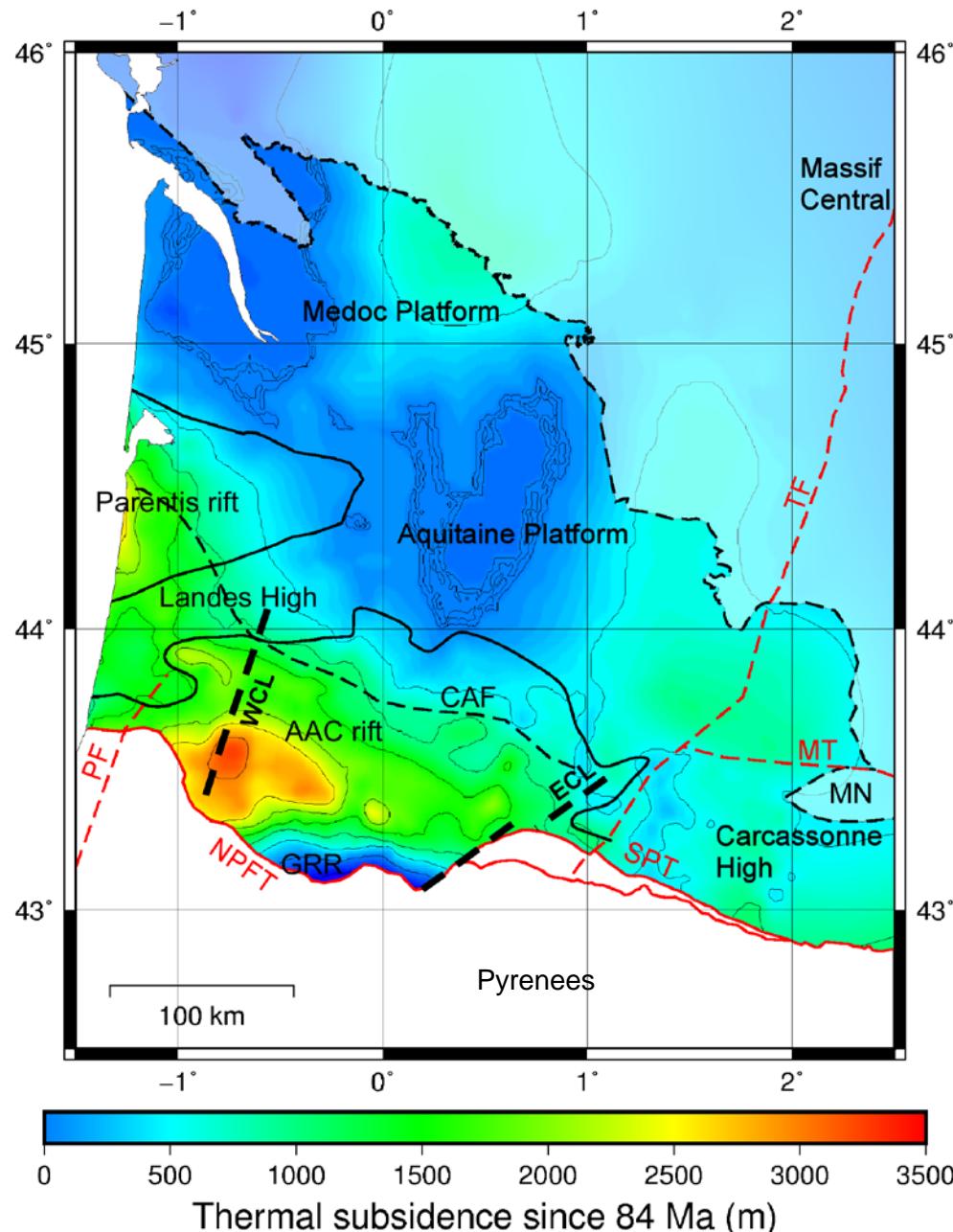
Angrand et al., in prep.



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Crustal template

Post-rift thermal subsidence



- Thermal subsidence during the Pyrenean orogeny (since 84 Ma)
- Uniform Stretching Model (McKenzie, 1978)
- Sediment-filled basin (2500 kg.m^{-3})
- Pre-rift lithospheric thickness: 125 km
- Thermal constant of the lithosphere: 50 Myr

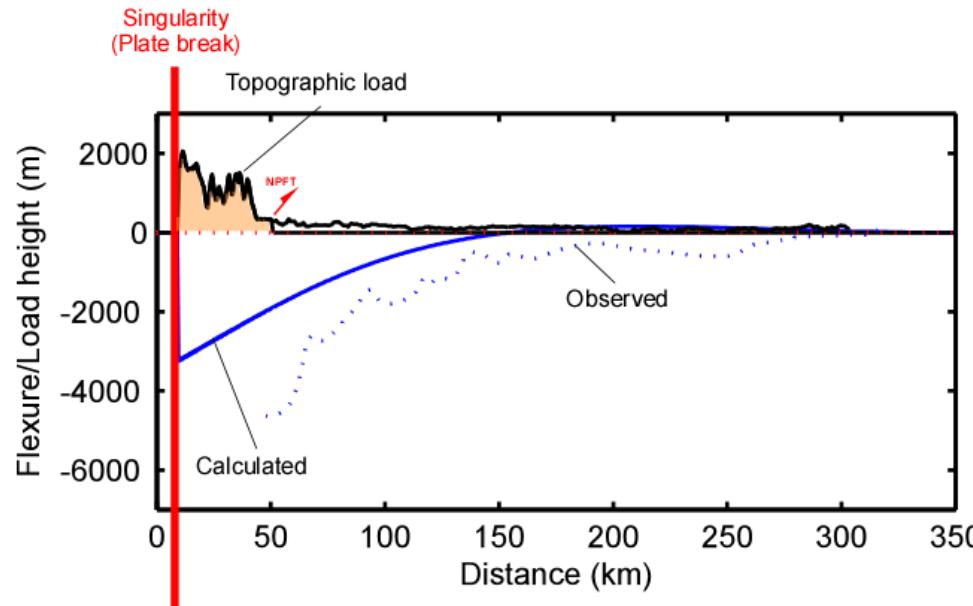
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Modelling the flexure

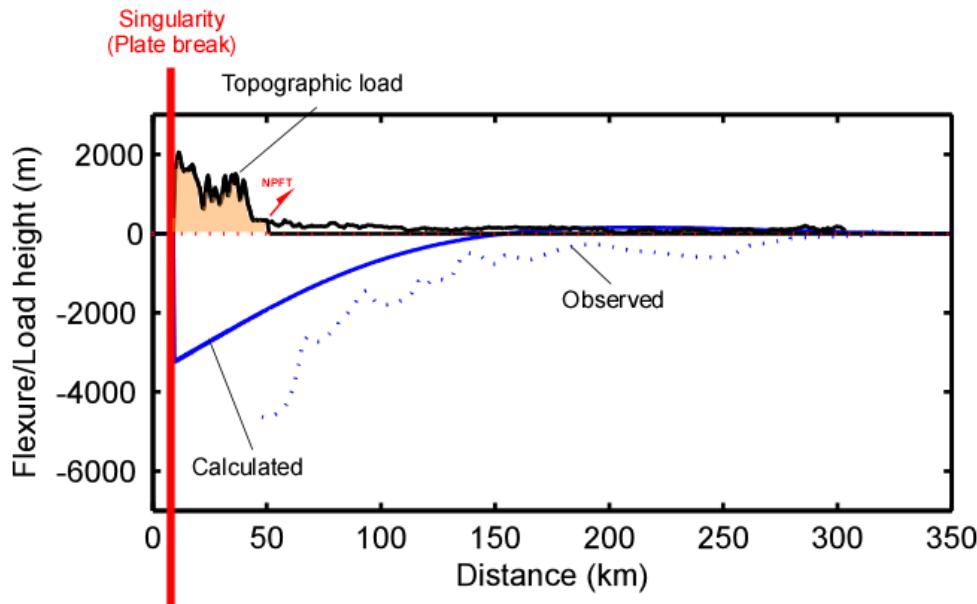
How it works



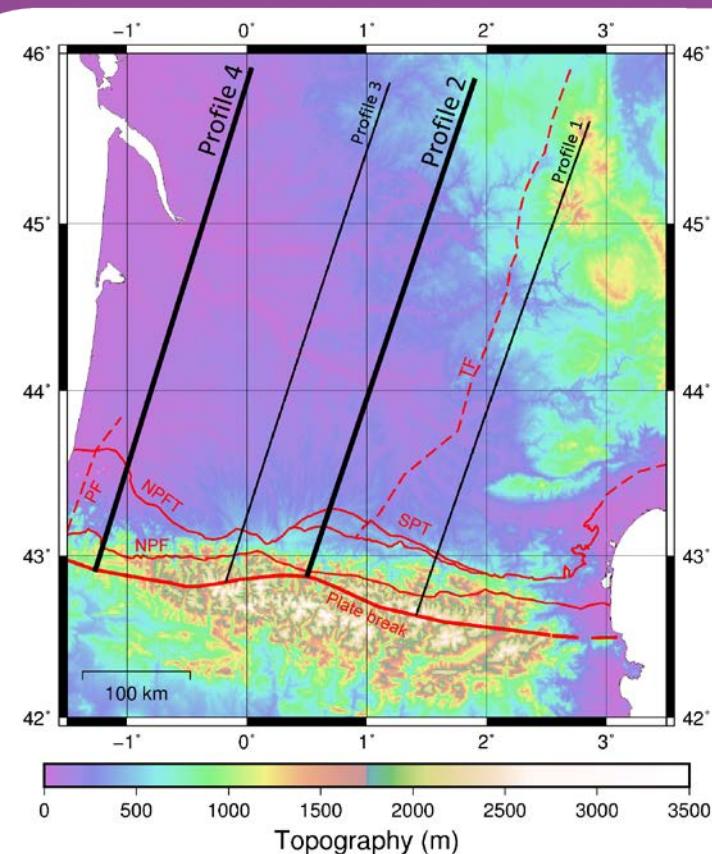
- Broken plate conditions
- Fixed physical parameters:
 - Young's modulus: 100 Gpa
 - Poisson's ratio: 0.25
 - Density of topographic load: 2700 kg.m^{-3}
 - Density of sediments: 2500 kg.m^{-3}
 - Density of mantle: 3200 kg.m^{-3}
- Varying T_e :
 - T_e is constant or varies along the profile
 - T_e varies along strike
- We compare the calculated flexure to the observed base of foreland surface

Modelling the flexure

How it works



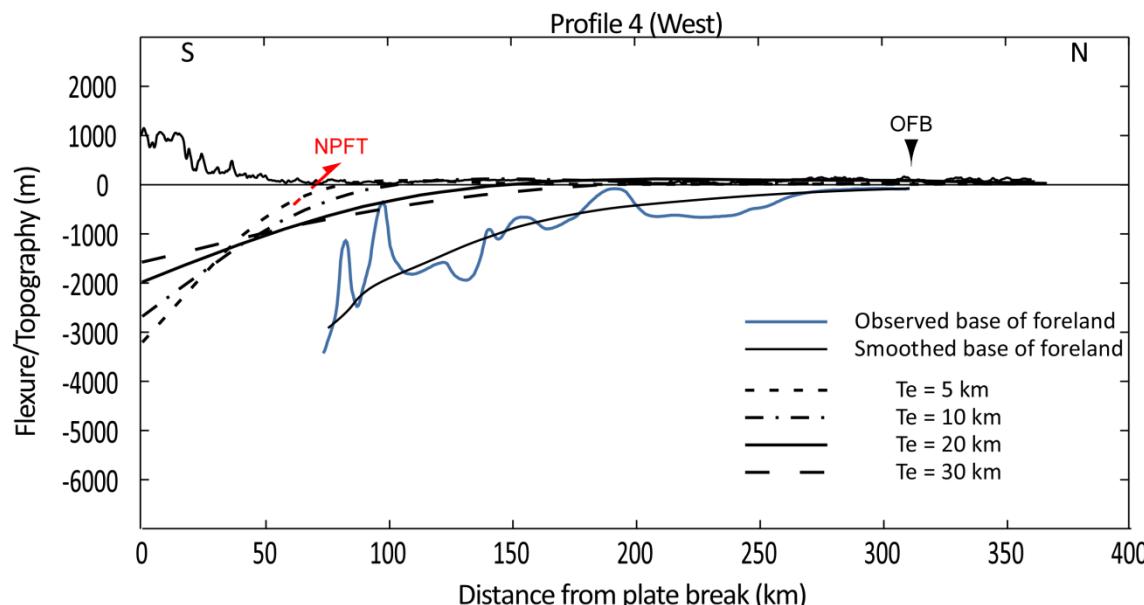
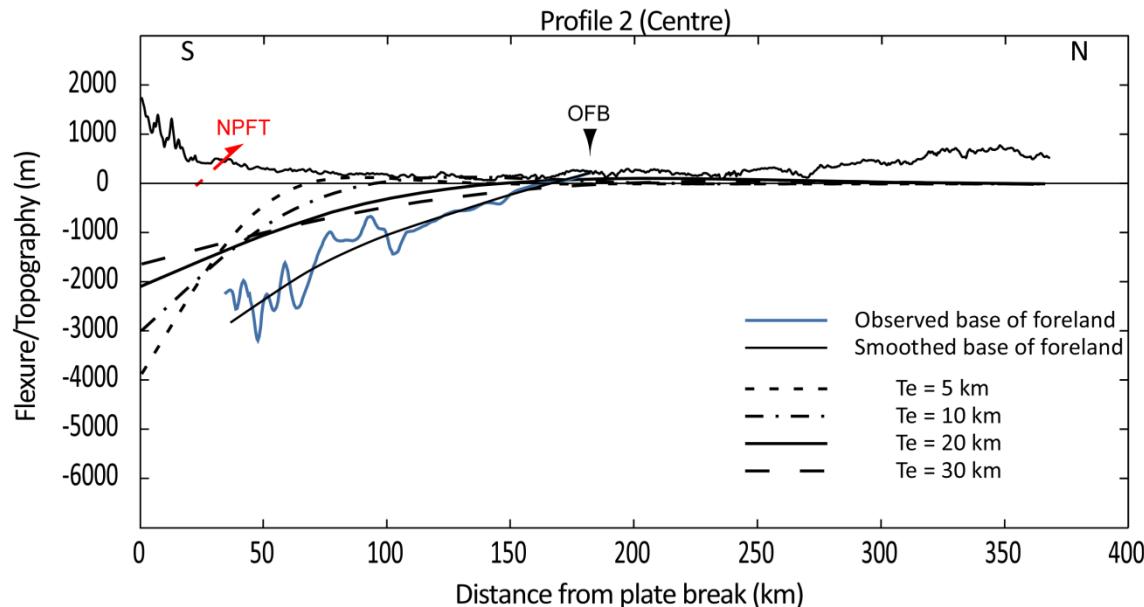
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Angrand et al., in prep.

Results

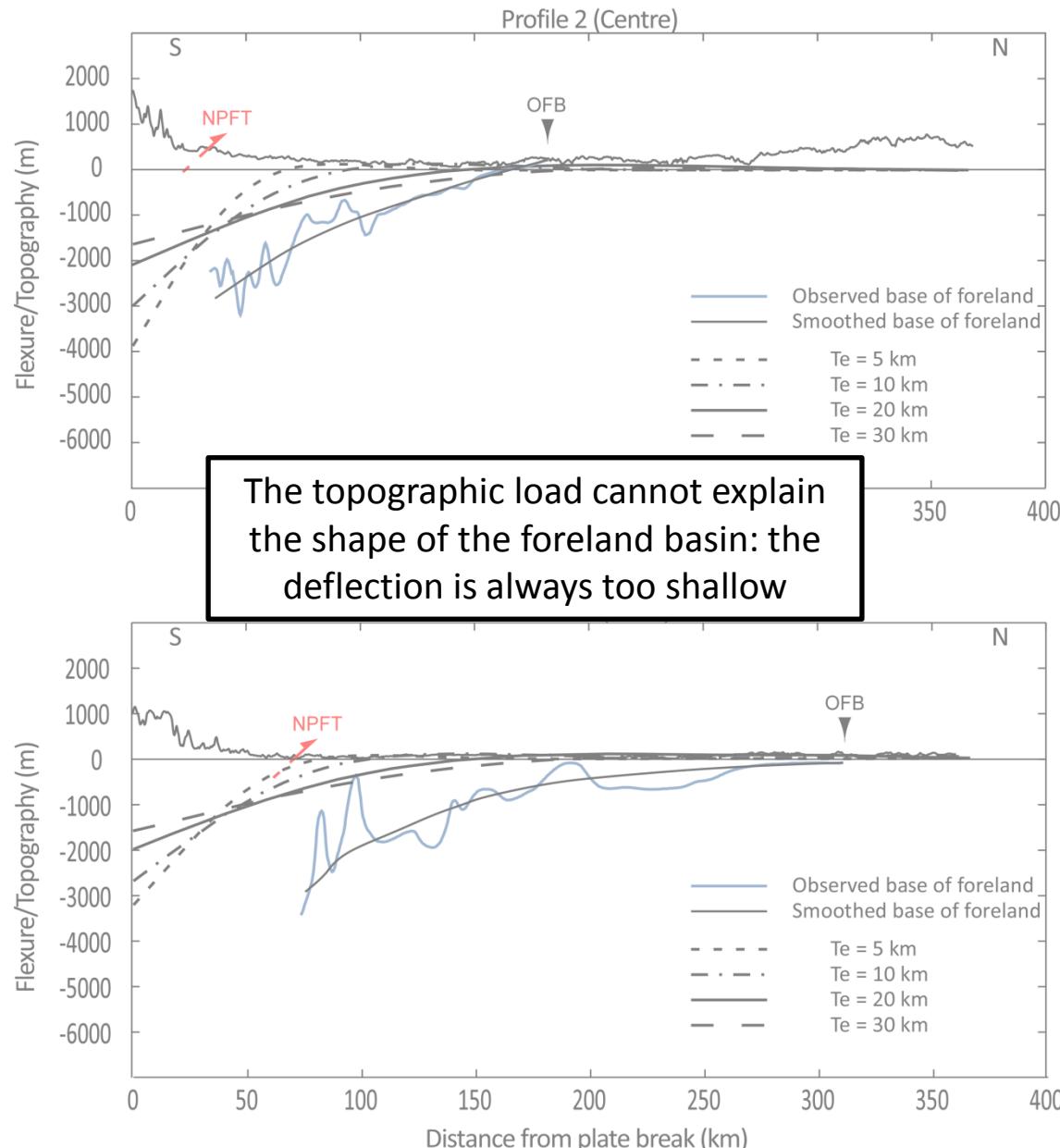
Flexure due to topography



Angrand et al., in prep.

Results

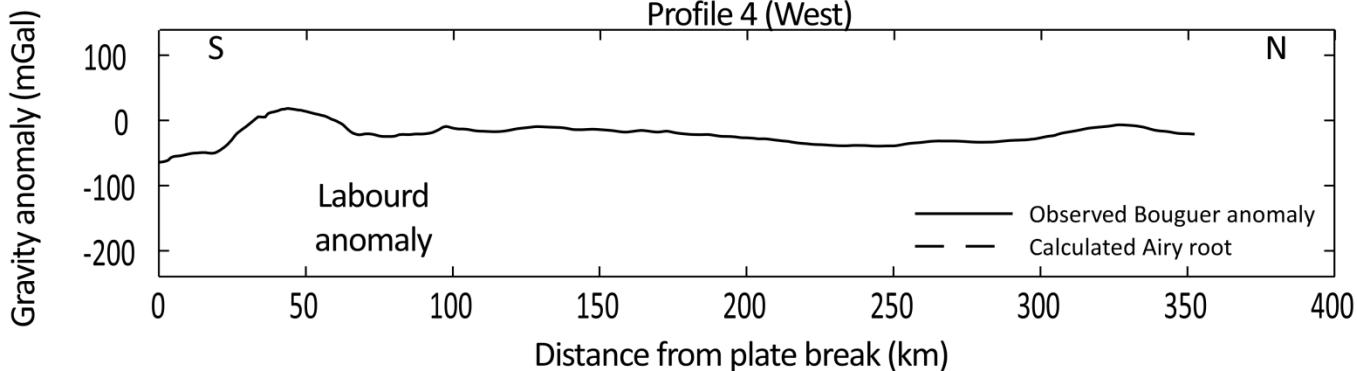
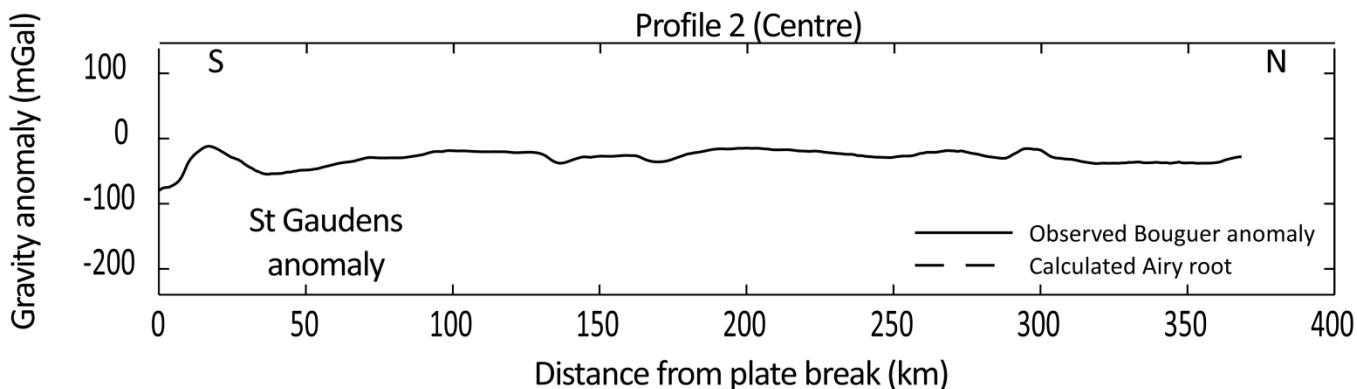
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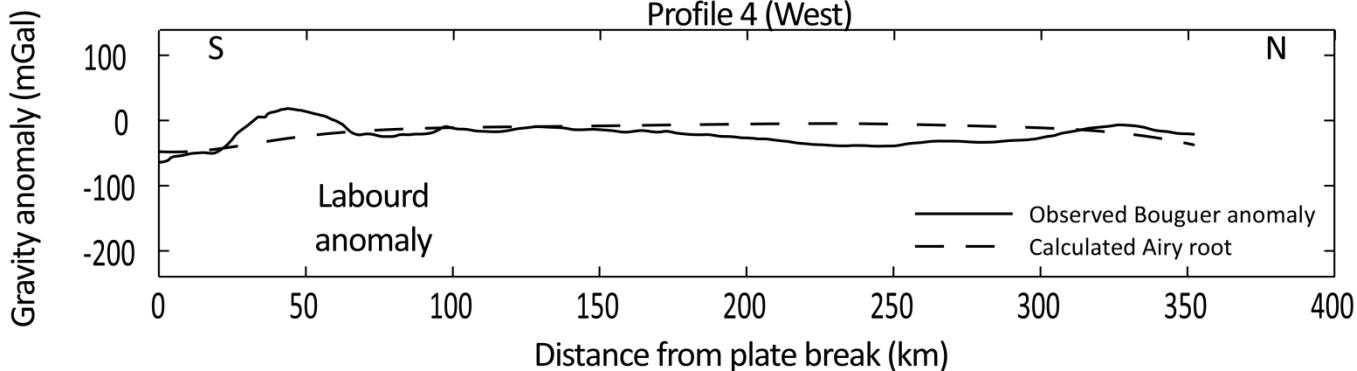
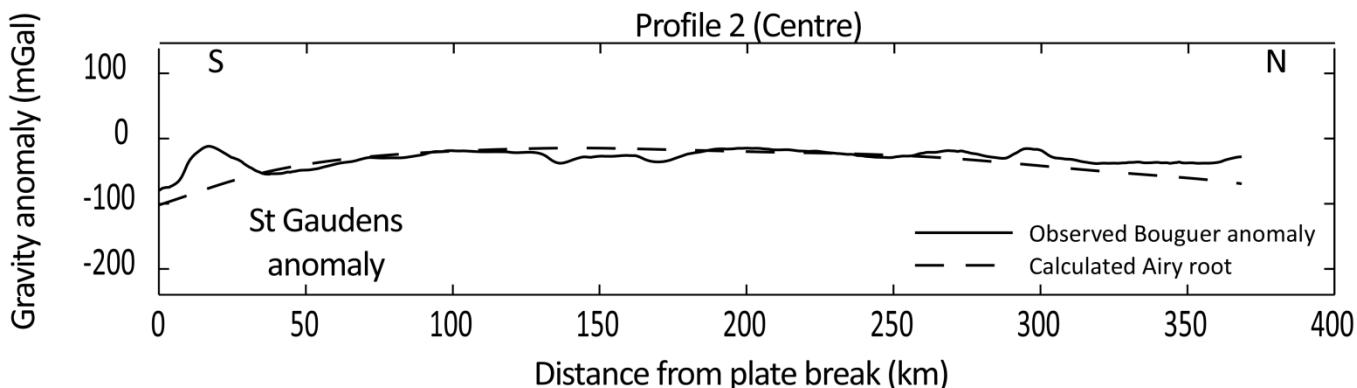
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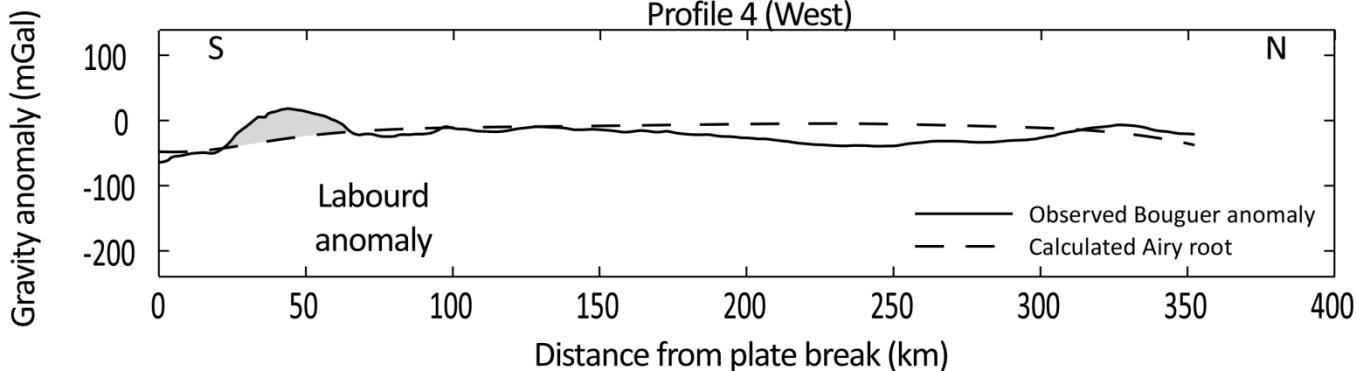
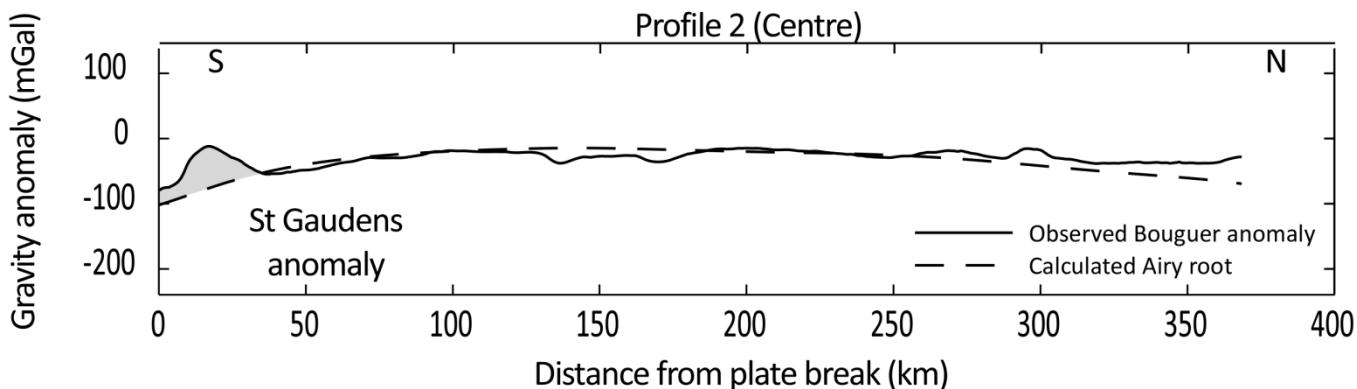
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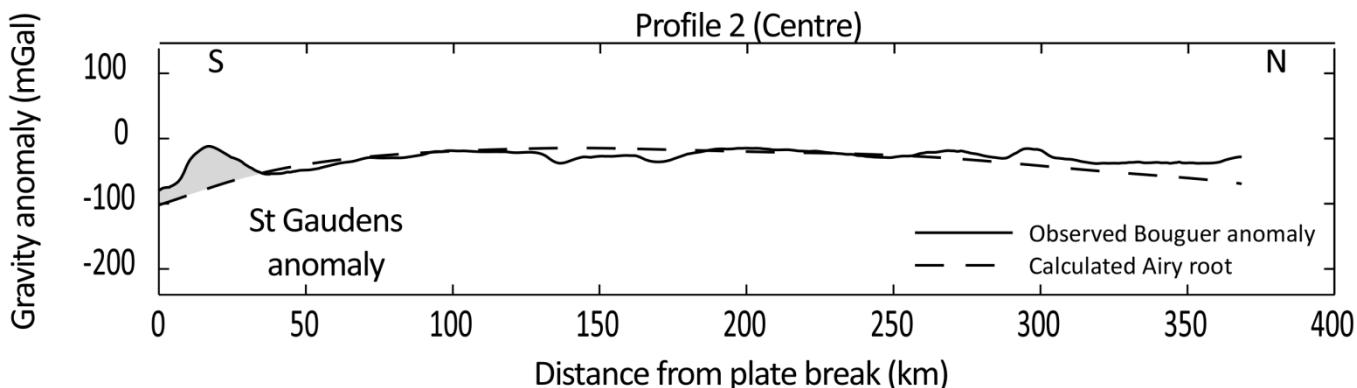
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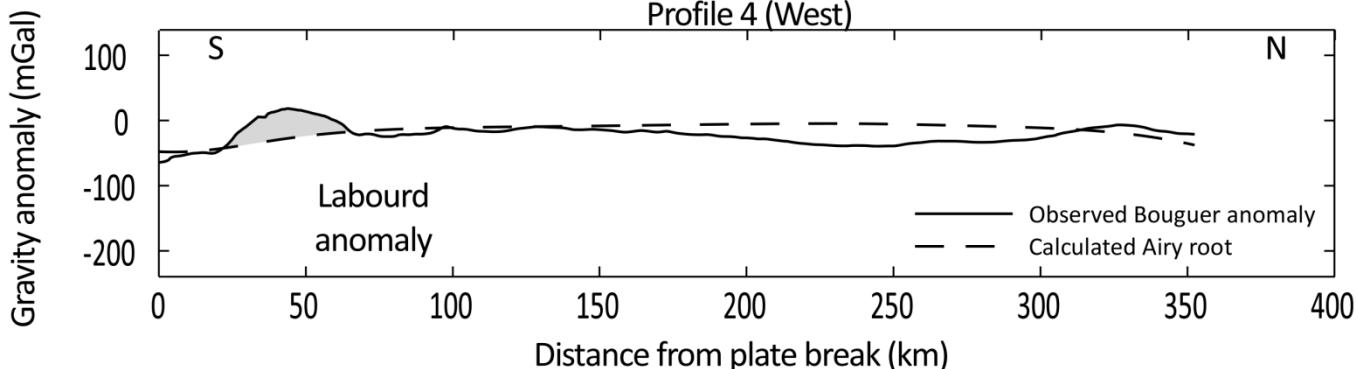
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Flexure due to topography



Green equivalent layer theorem:

$$H = \frac{\Delta g_{\text{Airy}}}{2\pi G \Delta \rho_b}$$

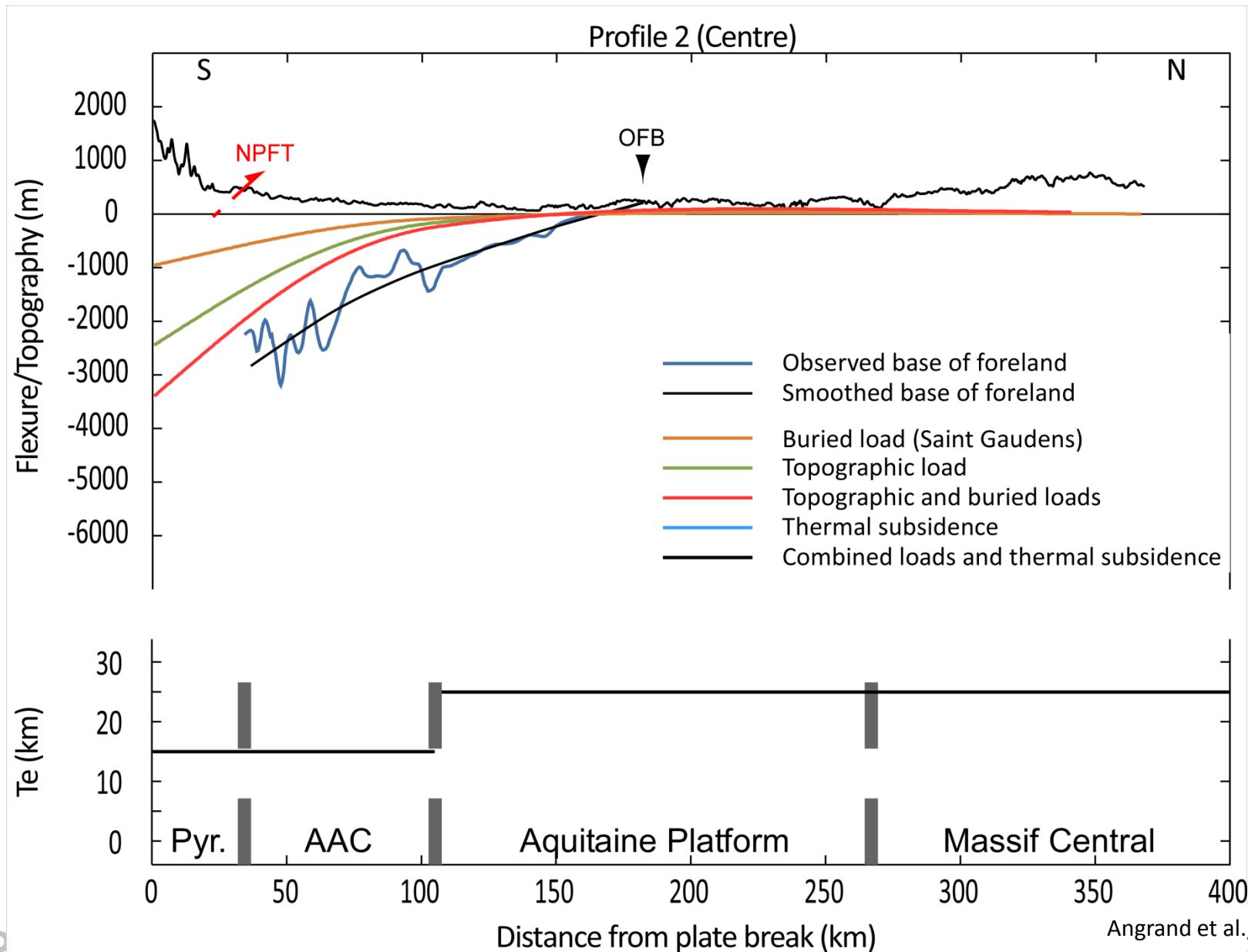


Δg_{Airy} : Positive part of The Airy isostatic anomaly

$\Delta \rho_b$: Density contrast:
500 kg.m⁻³

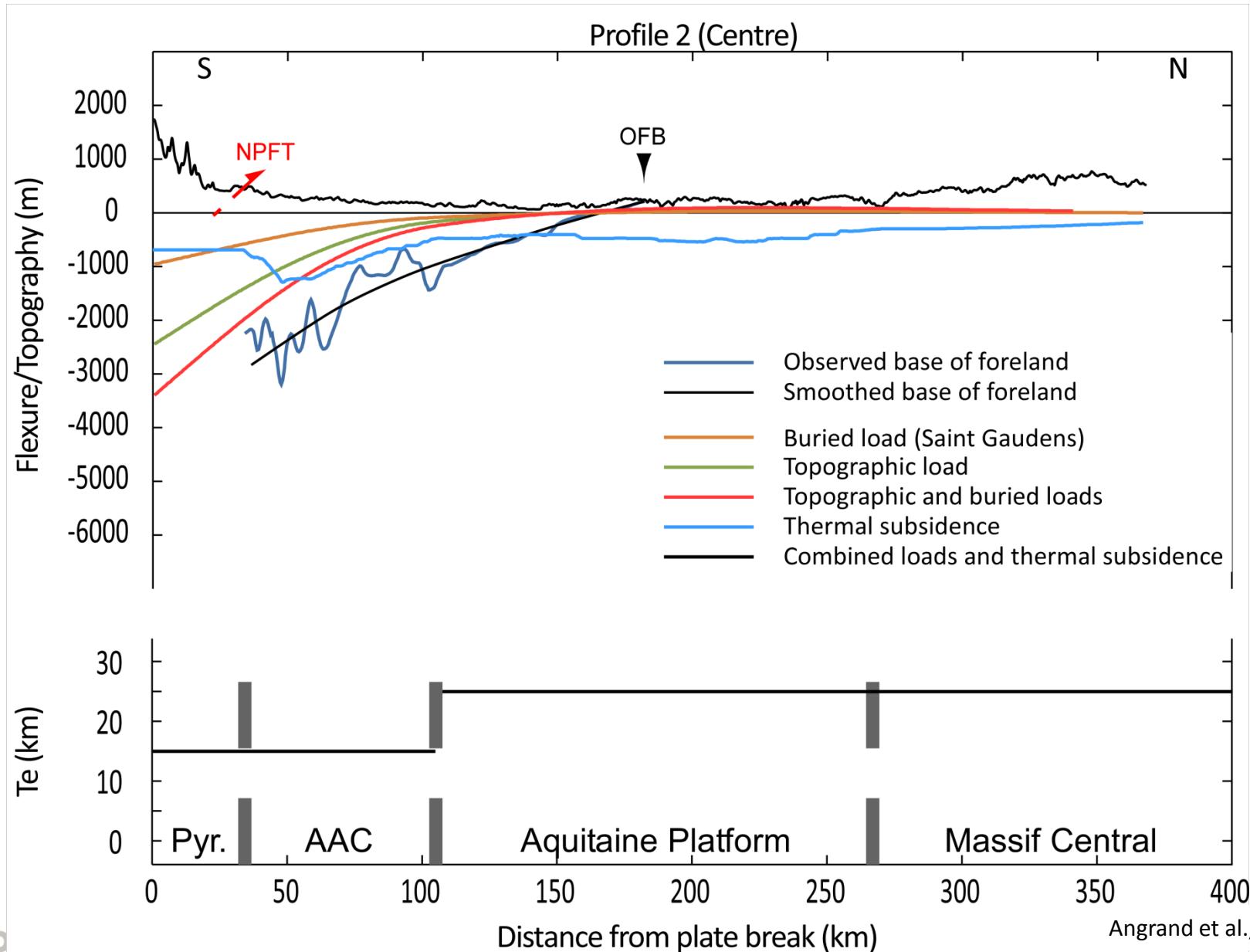
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Combined subsidence: Central Aquitaine Basin



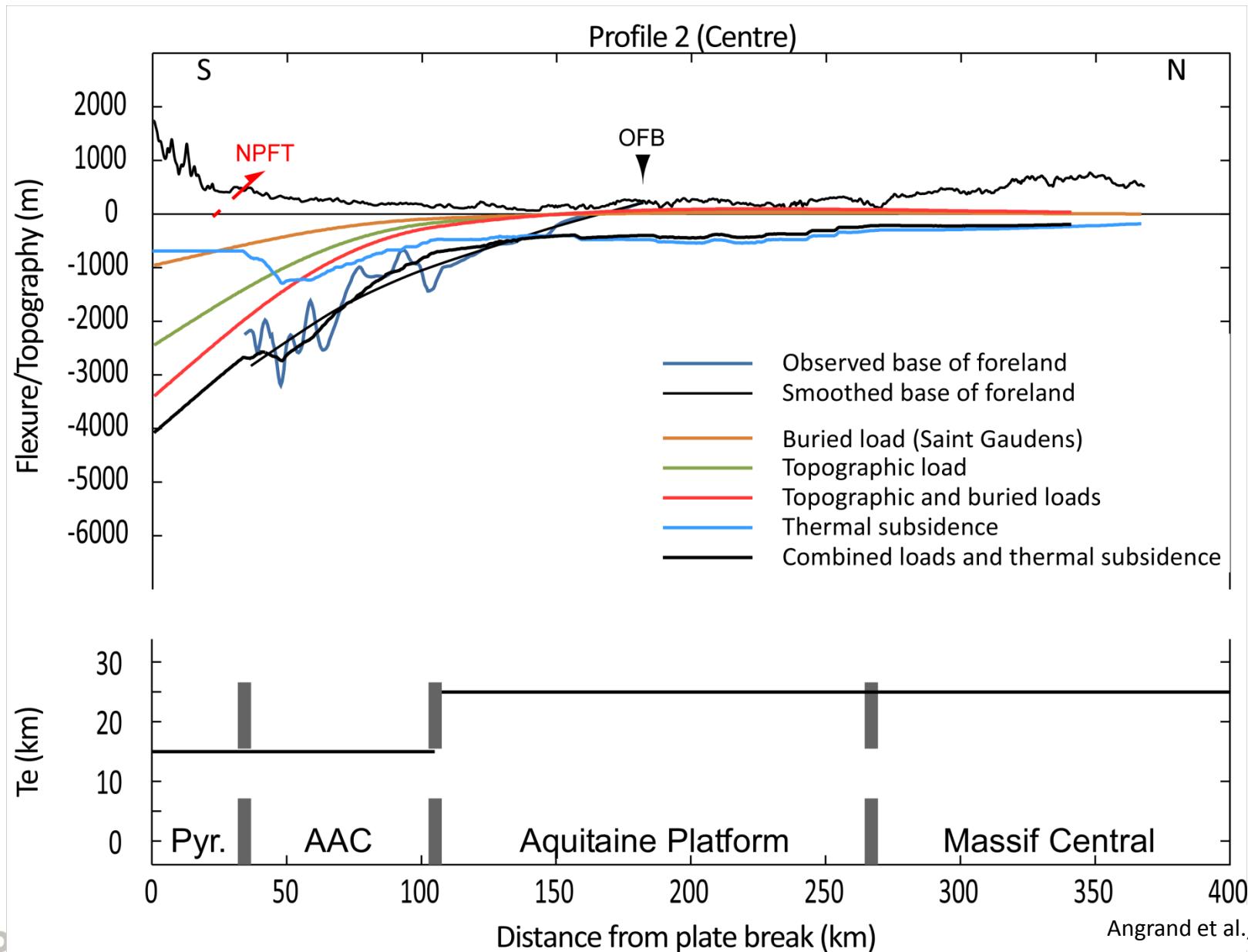
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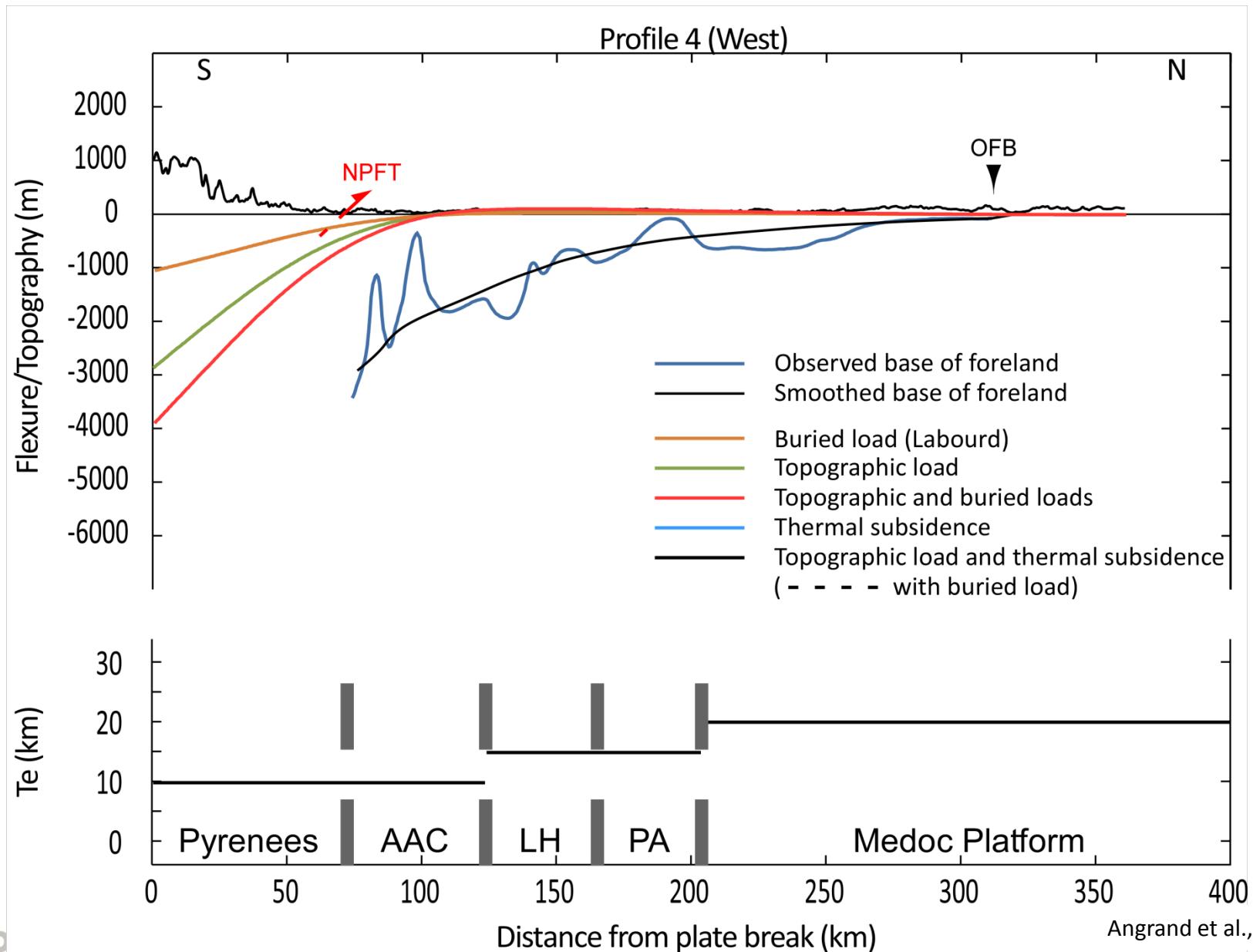
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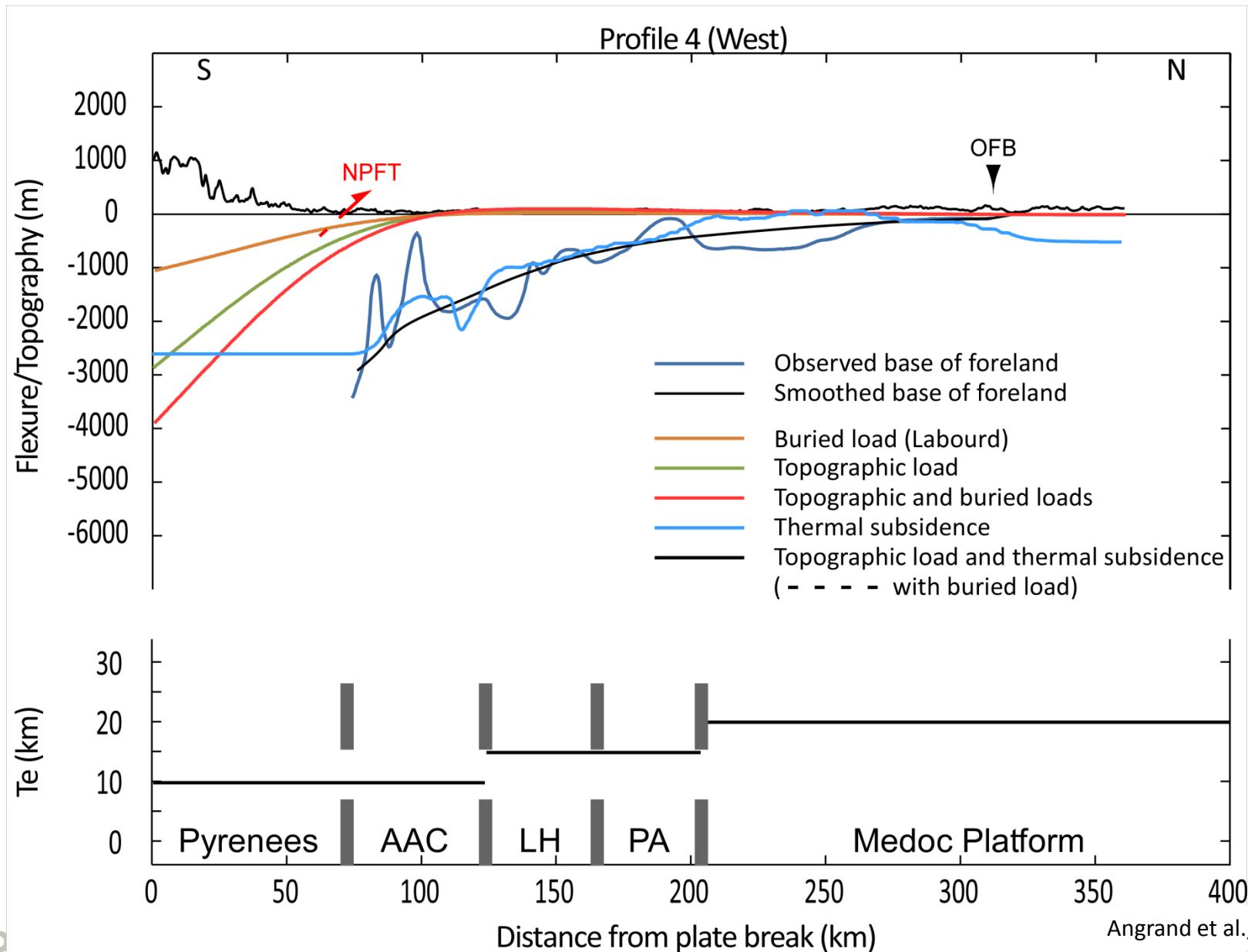
Results

Combined subsidence: Western Aquitaine Basin



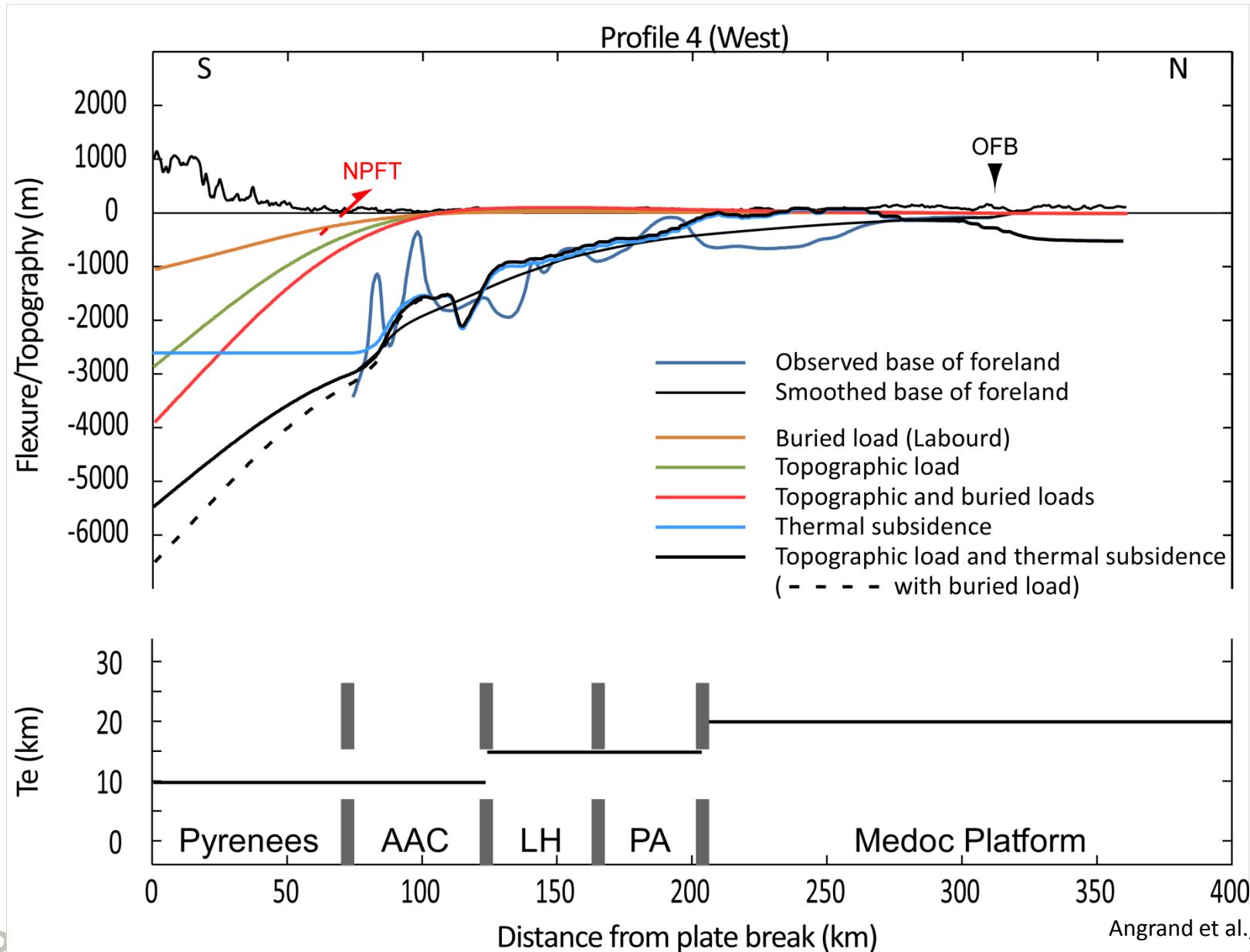
Results

Combined subsidence: Western Aquitaine Basin



Results

Combined subsidence: Western Aquitaine Basin



- The Apto-Cenomanian rifting affected the evolution of the Aquitaine foreland basin
 - Strong segmentation of both basement and foreland
 - Post-rift thermal subsidence was active during the Pyrenean orogeny
 - T_e decreases from centre to west and from north to south (from 7 to 25 km), consistent with variations of β

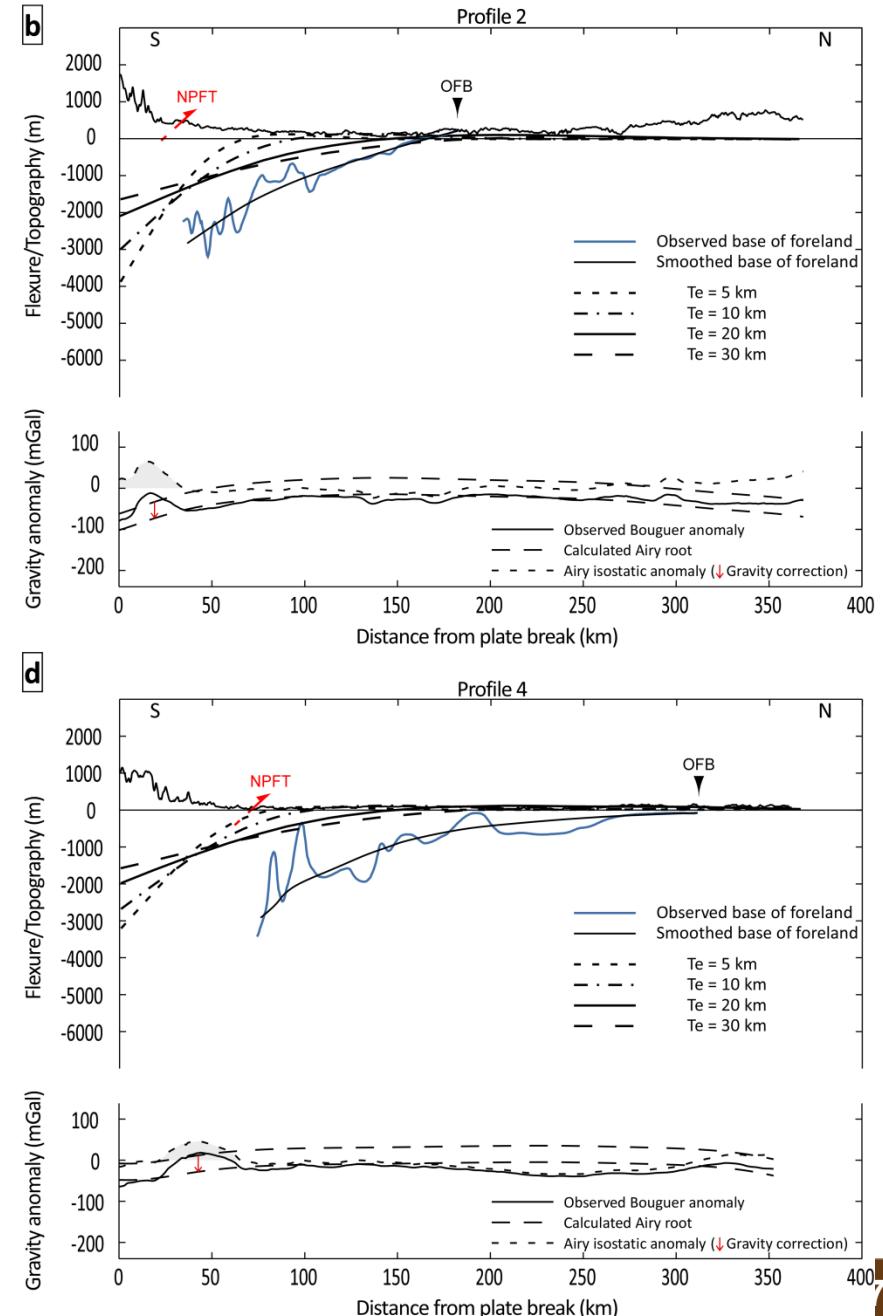
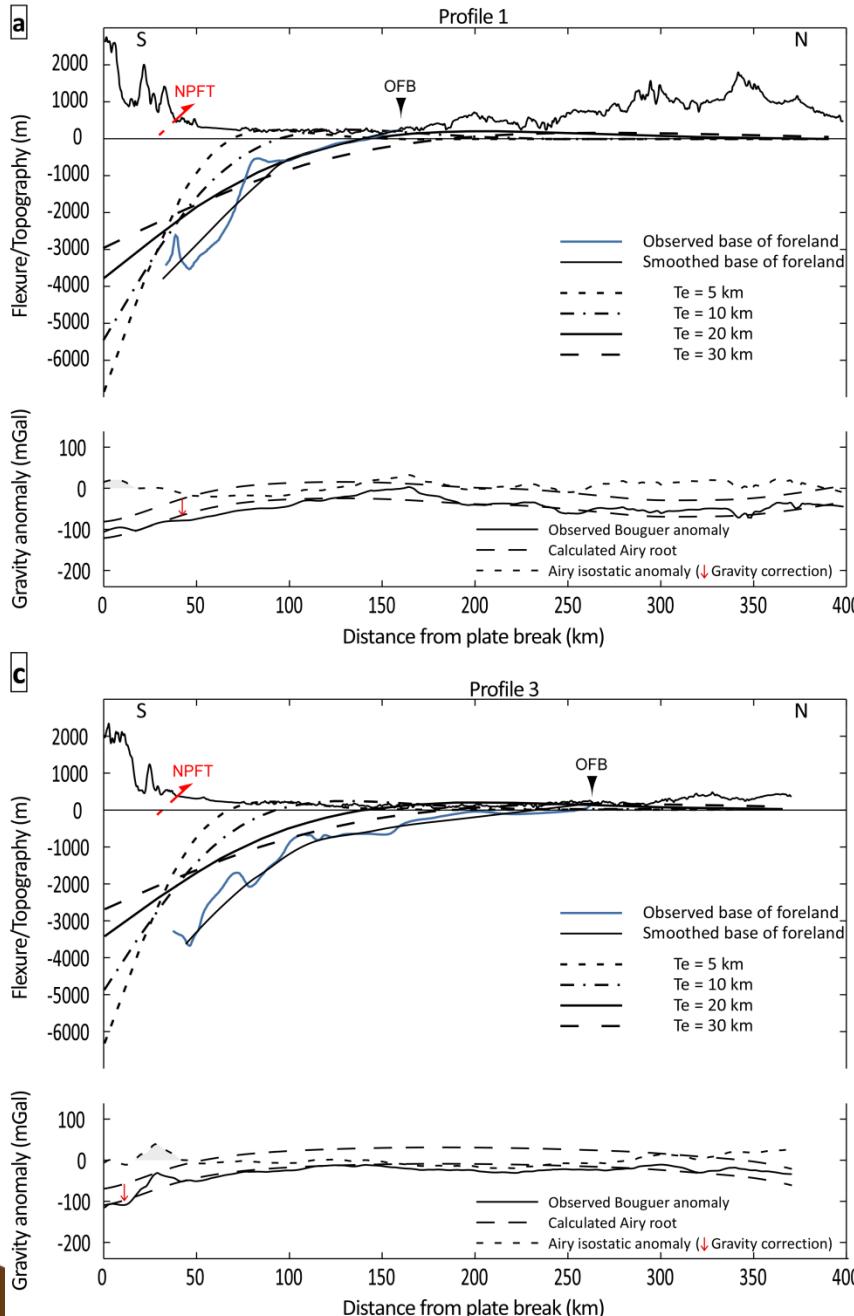
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 - What about the Labourd gravity anomaly?

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 - Pyrenean topography cannot explain the deflection of the European plate
 - The geometry of the foreland base of the central Aquitaine Basin is explained by combined loading (topographic and buried loads) and thermal subsidence
 - In the westernmost Aquitaine Basin, the foreland basin geometry is mainly controlled by thermal subsidence

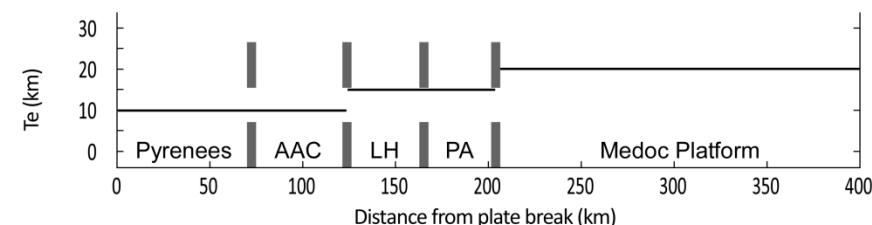
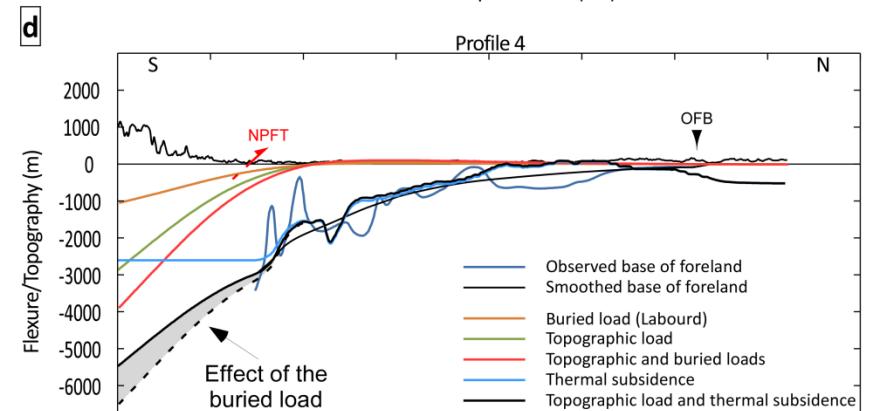
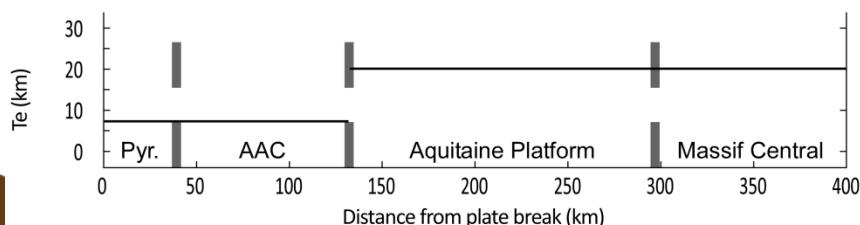
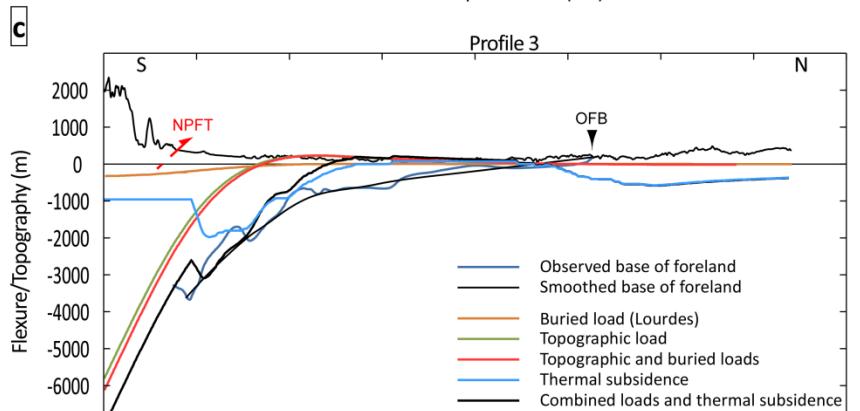
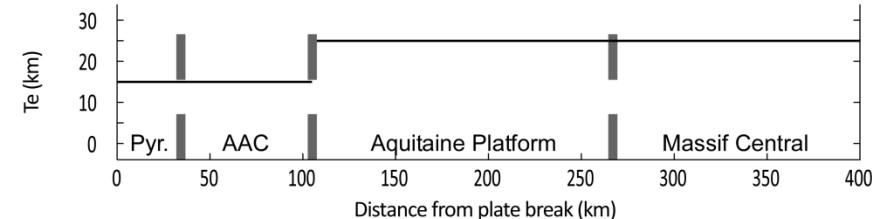
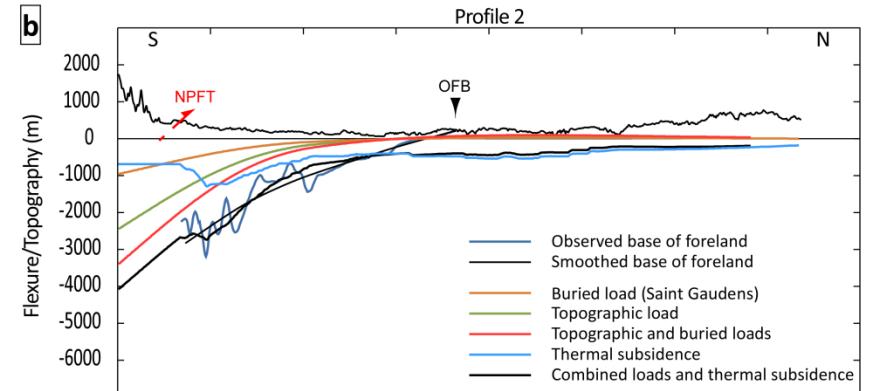
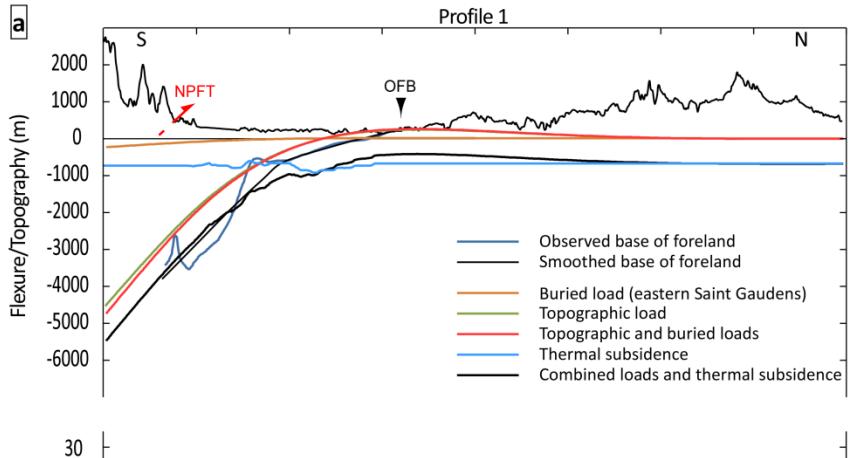
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- Controlling factors of the flexure
 - Loading of a rifted crust affects the development of flexural foreland basin (if lithosphere is not thermally equilibrated)
 - In this case, thermal cooling can contribute significantly to subsidence
 - Qualitative correlation between β and T_e : the lower the T_e , the higher the β

Thanks for your attention

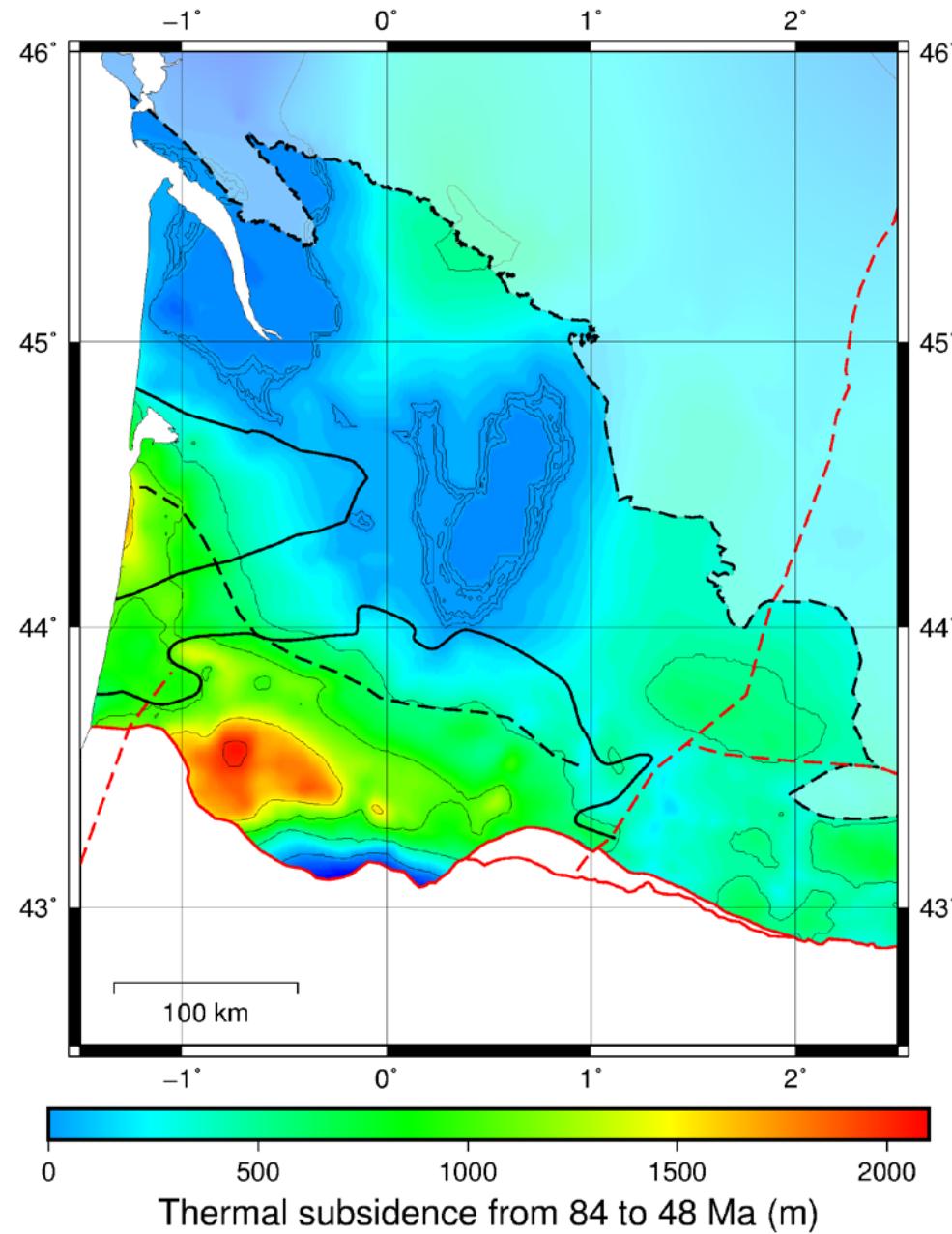
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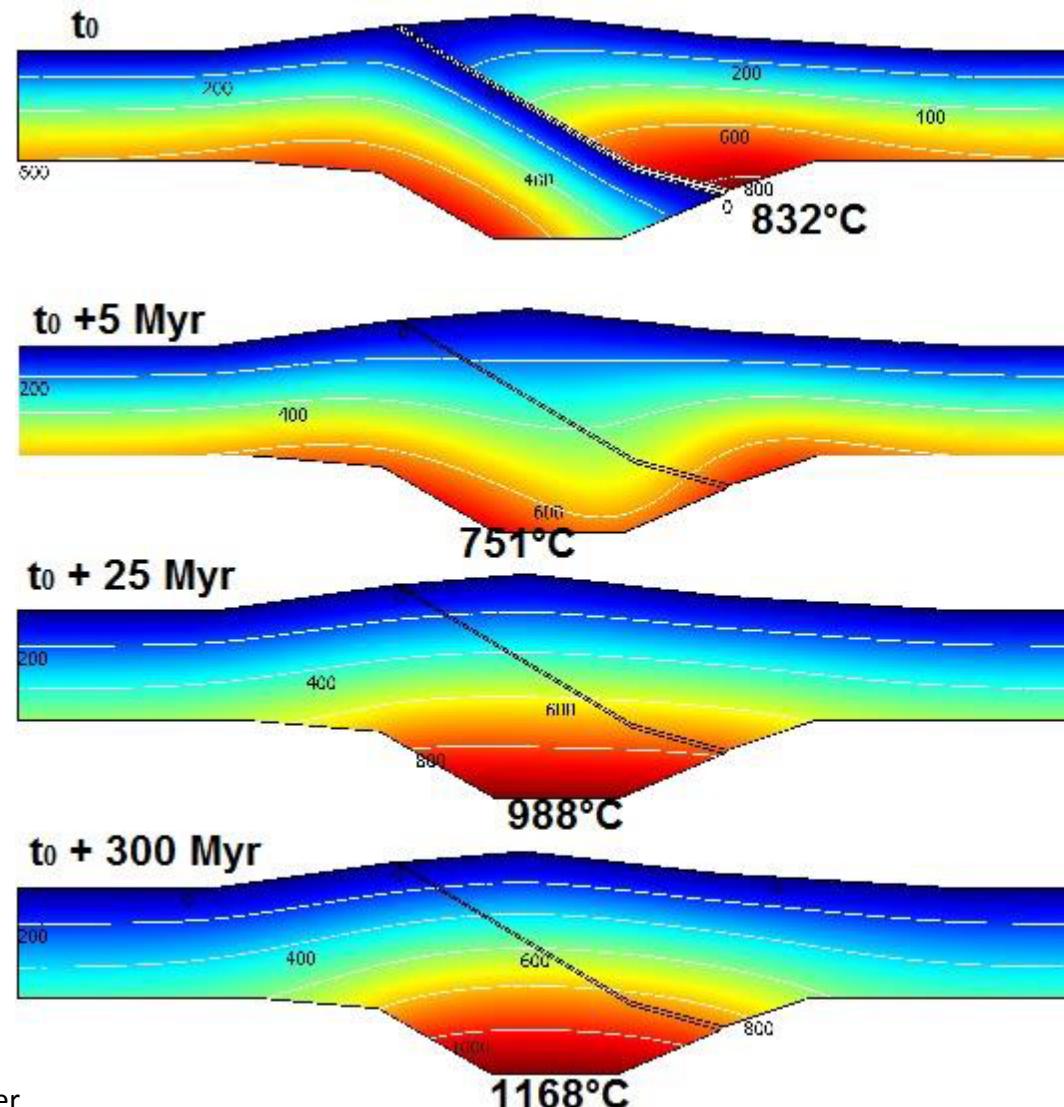
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++ | Thermal subsidence



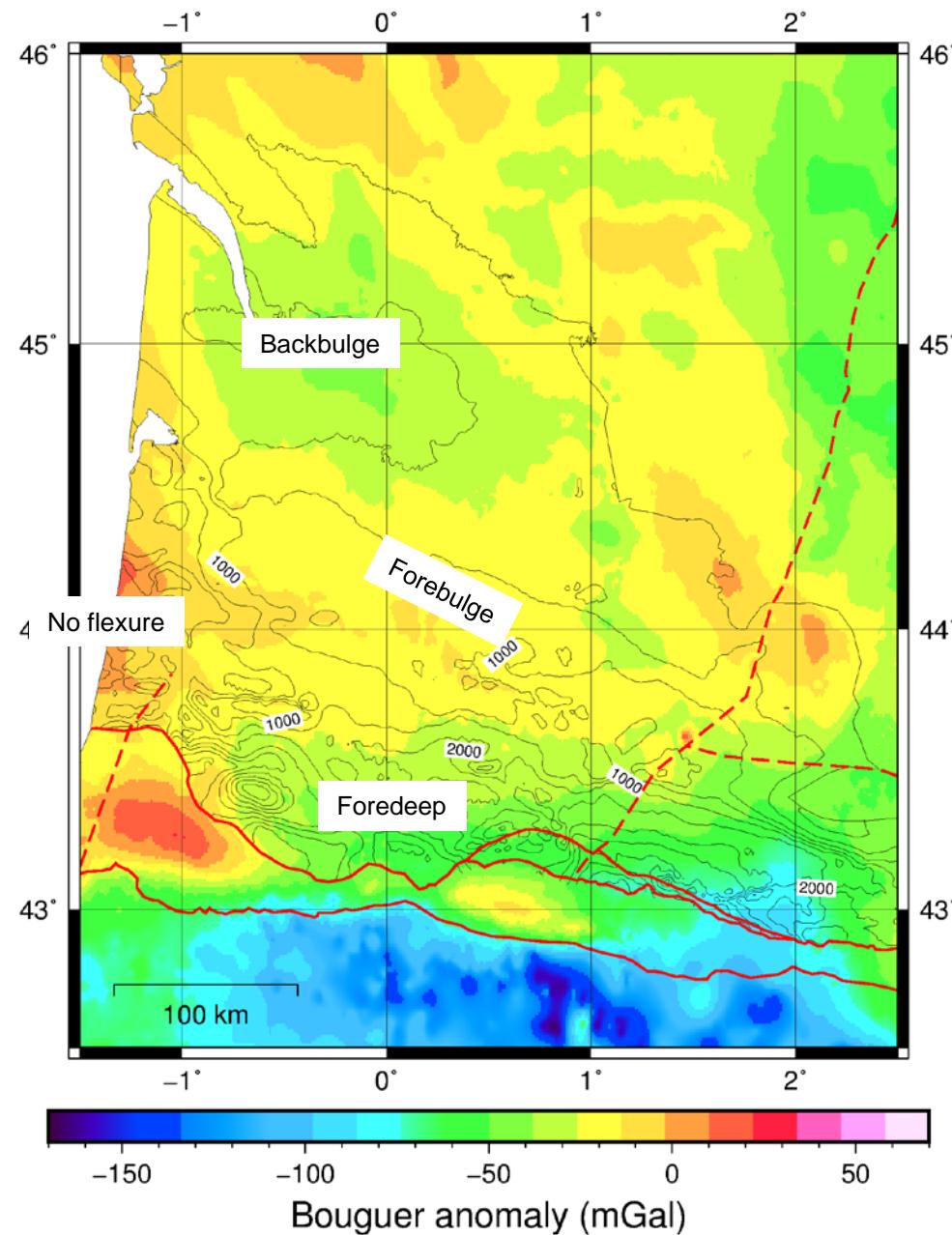
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Courtesy of L. Guillou-Frottier

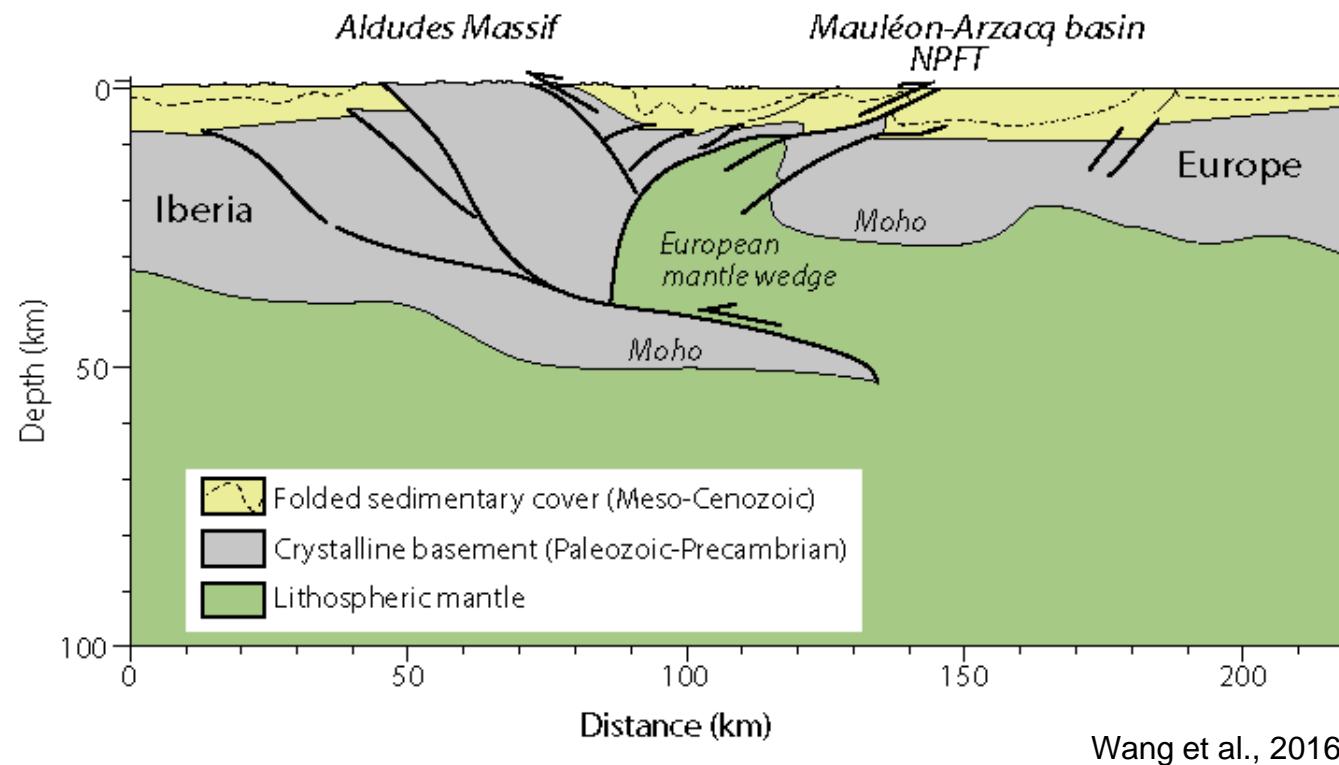
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Gravity (Bouguer anomaly)

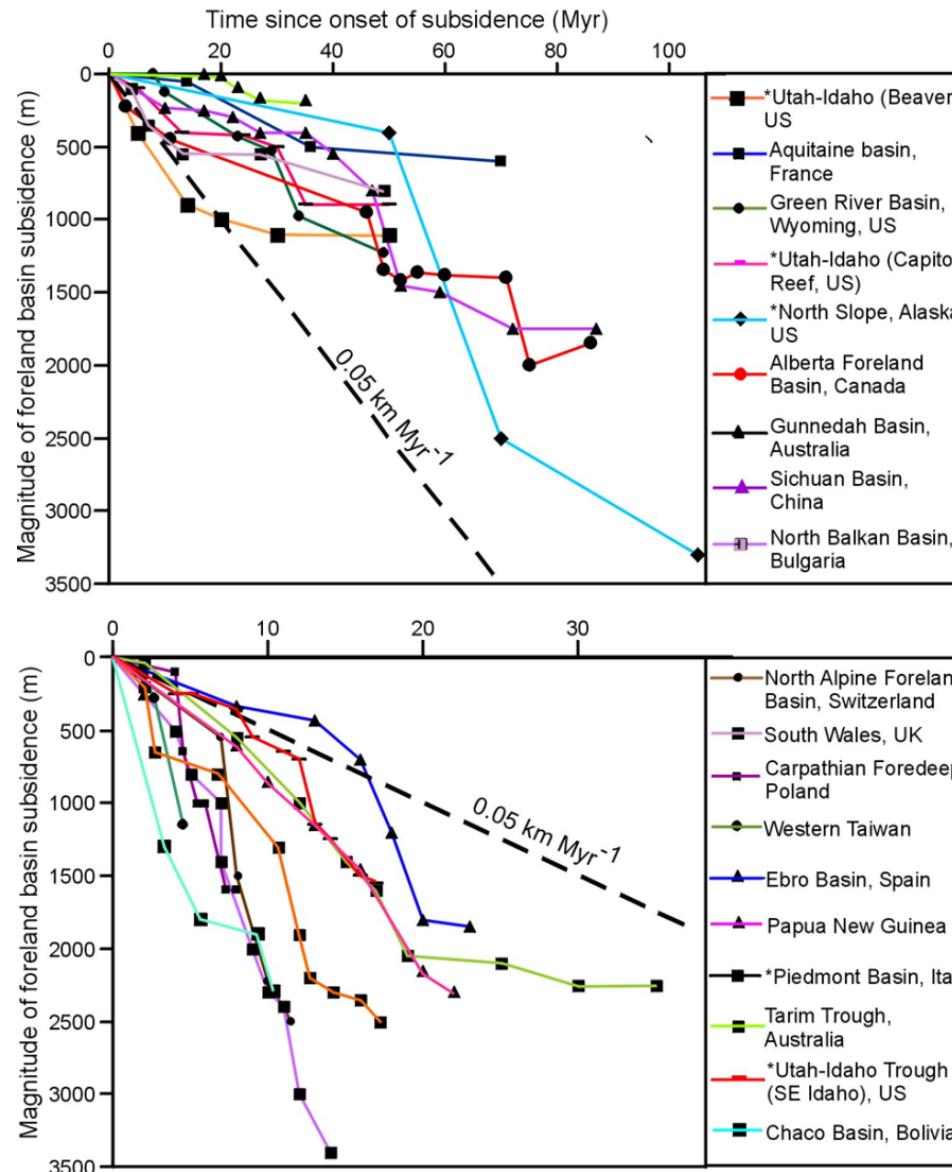


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No flexure in the west?

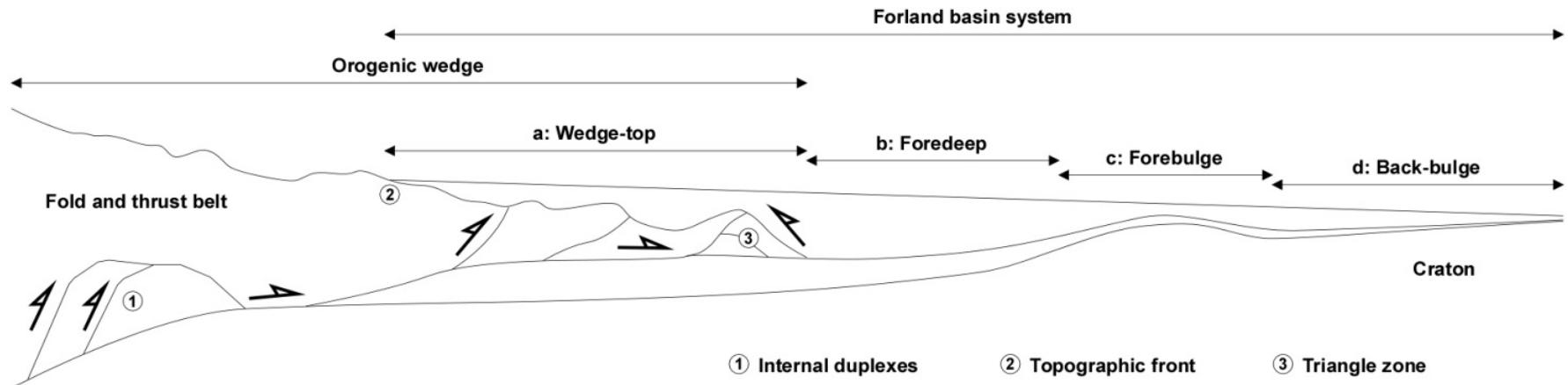


++ Subsidence in foreland basins



Naylor and Sinclair, 2008

++ | Foreland basins



DeCelles and Giles, 1996

++ | Gravity in foreland basin

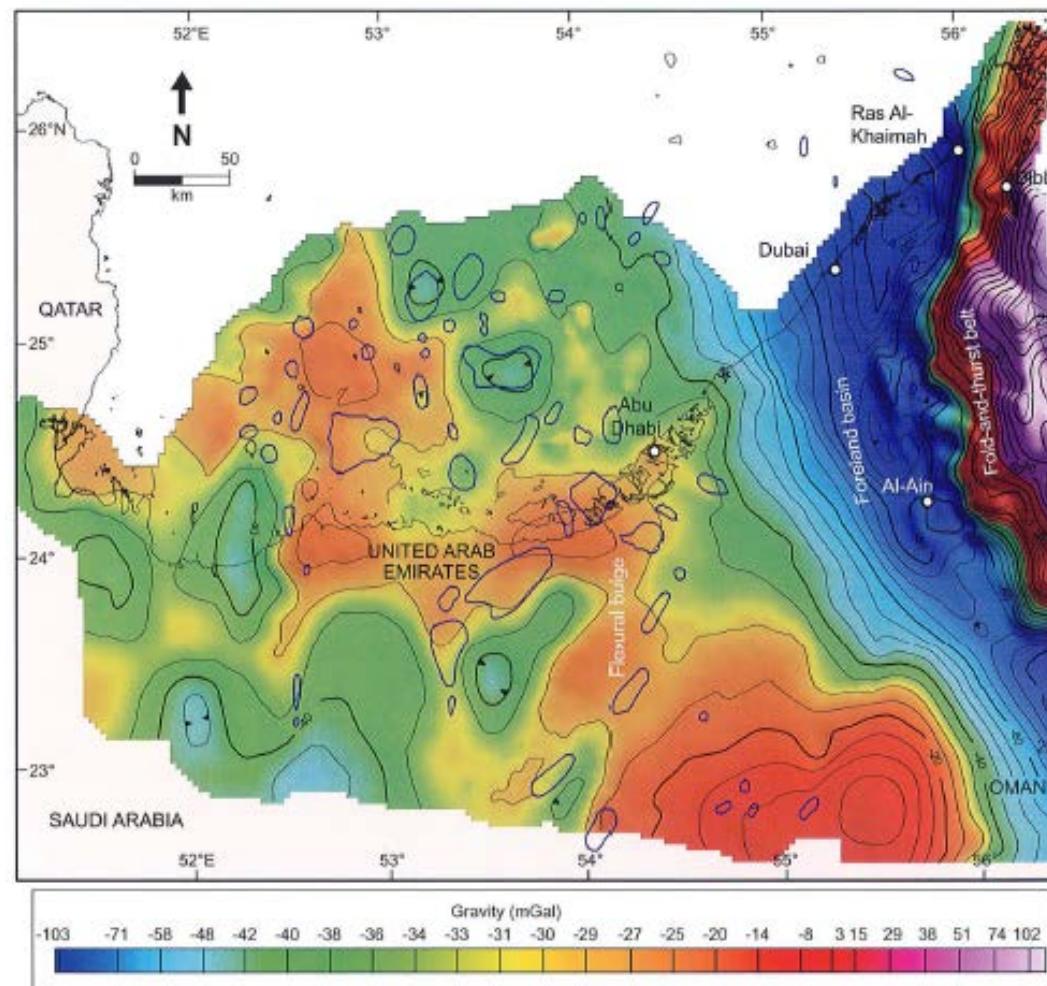
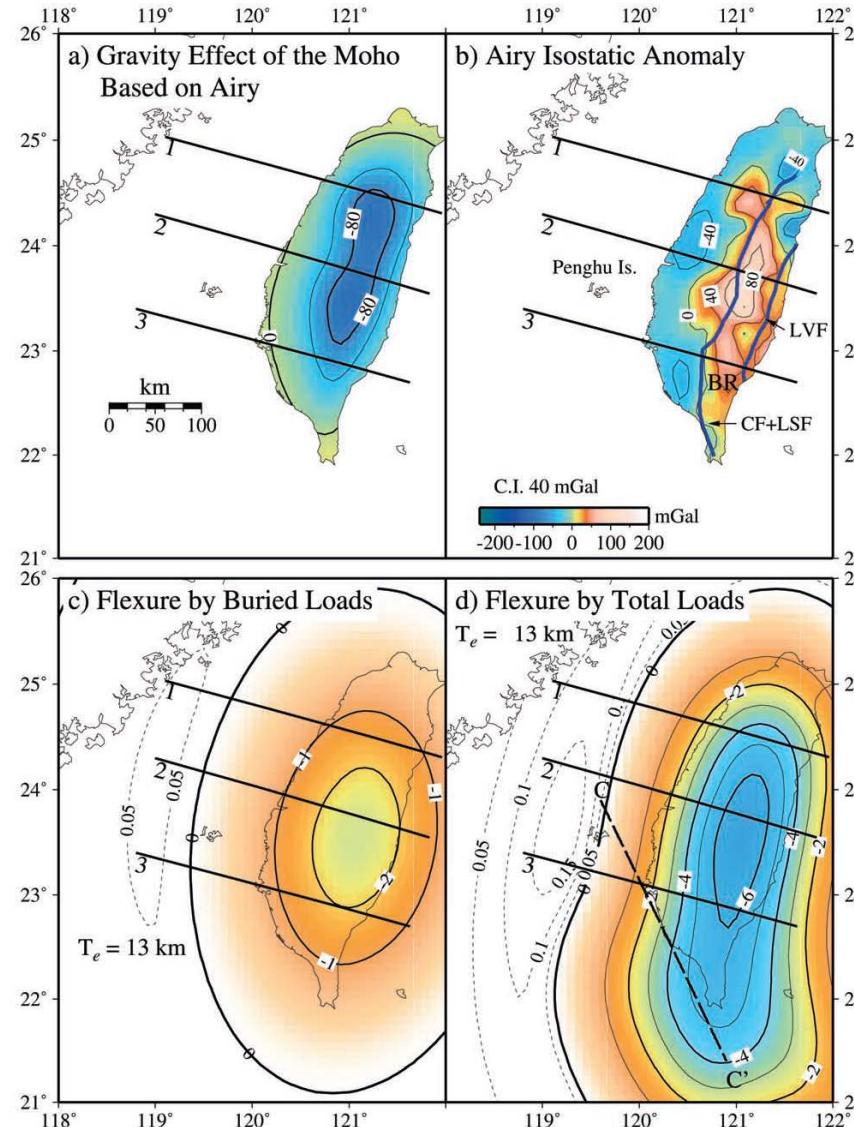


Figure 4: Bouguer anomaly map of the UAE and surrounding areas. Illumination from the northeast. The blue outlines show location of oil and gas fields. See Figure 2 for names of fields. See Enclosure I for enlarged version.

Ali et al., 2014

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Flexure



Lin and Watts, 2002

